

TEMPEST

TESTING AND EVALUATING MODEL PREDICTIONS
OF EUROPEAN STORMS

Large-scale temperature gradients and the extratropical storm tracks

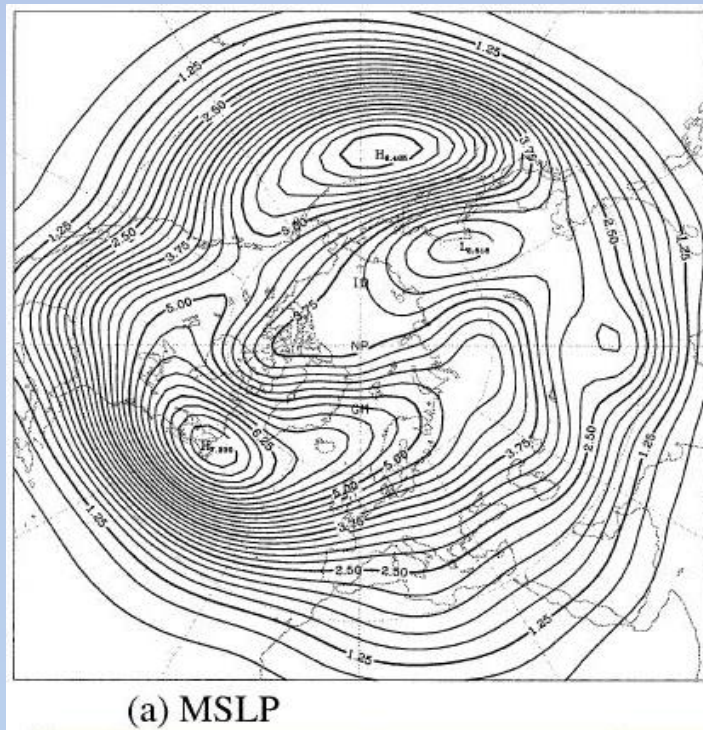
Ben Harvey, Len Shaffrey, Tim Woollings
NCAS Climate/University of Reading

CWC/SRM Annual meeting, Oxford, 14th May 2013

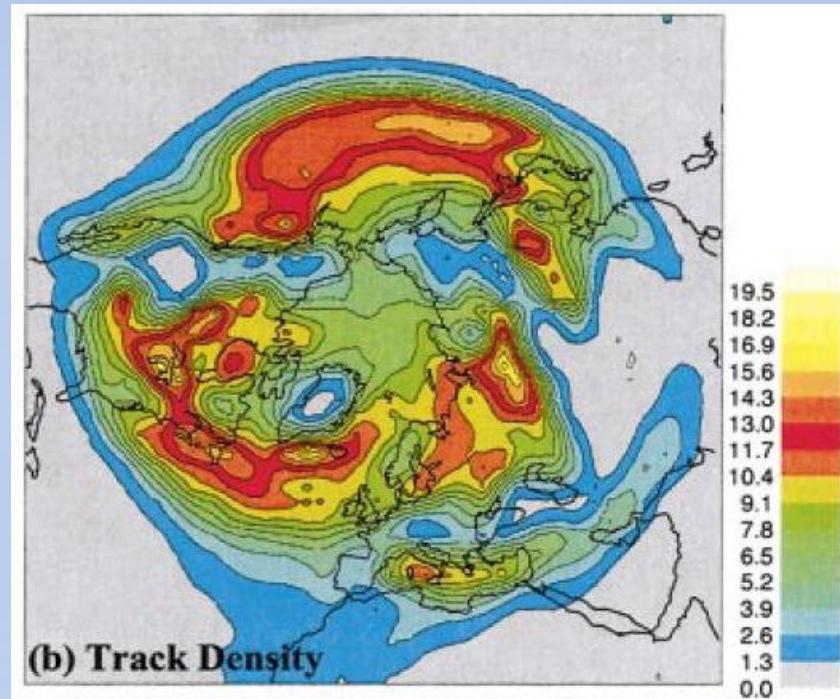
Lots of storm track diagnostics – which to use?

- Depends on purpose and data availability
- ‘Eulerian’ or ‘feature tracking’?

2-6 day MSLP standard deviation



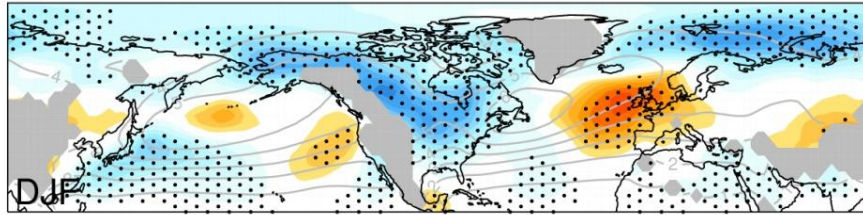
Feature tracking using 850 hPa vorticity



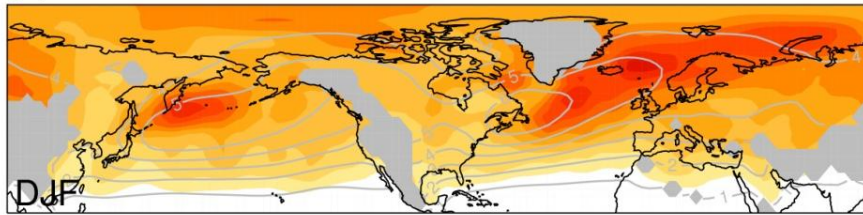
Hoskins and Hodges (2002)

Storm track responses in CMIP5

(a) winter multi-model mean



(b) winter inter-model std dev



Data from 24 models
One run per model

Scenario:
RCP8.5 – HISTORICAL

Storm track measure:
2-6 day MSLP std dev

Possible causes of the spread

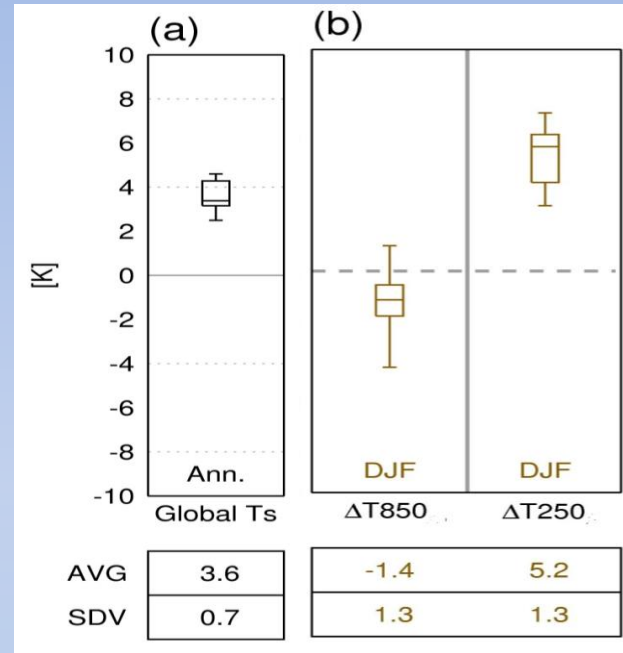
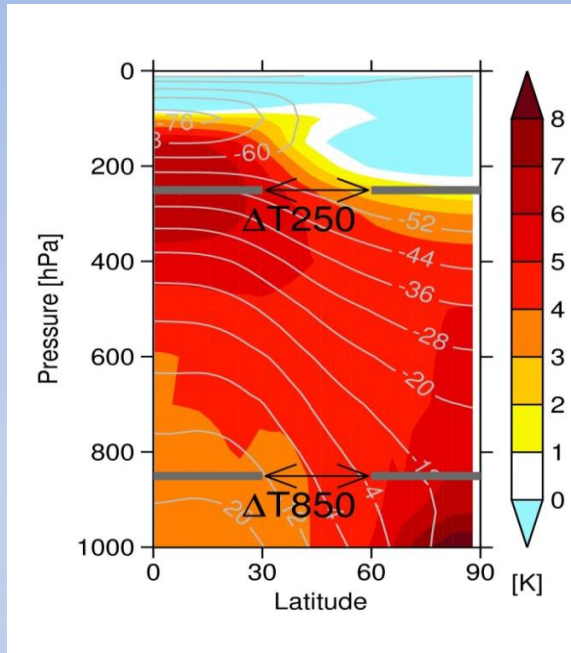
Natural variability

Baroclinicity (horizontal T gradients vs static stability)

Diabatic processes

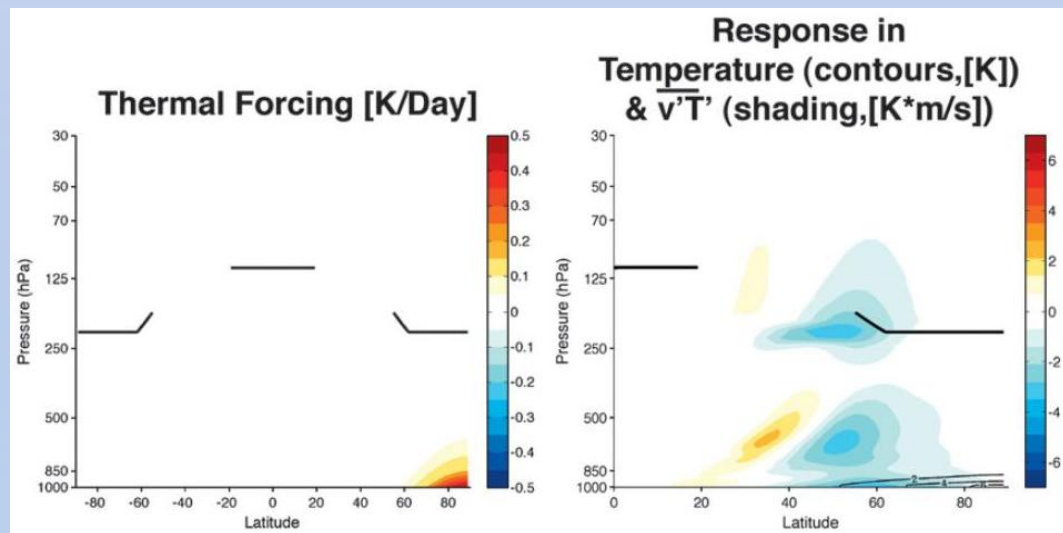
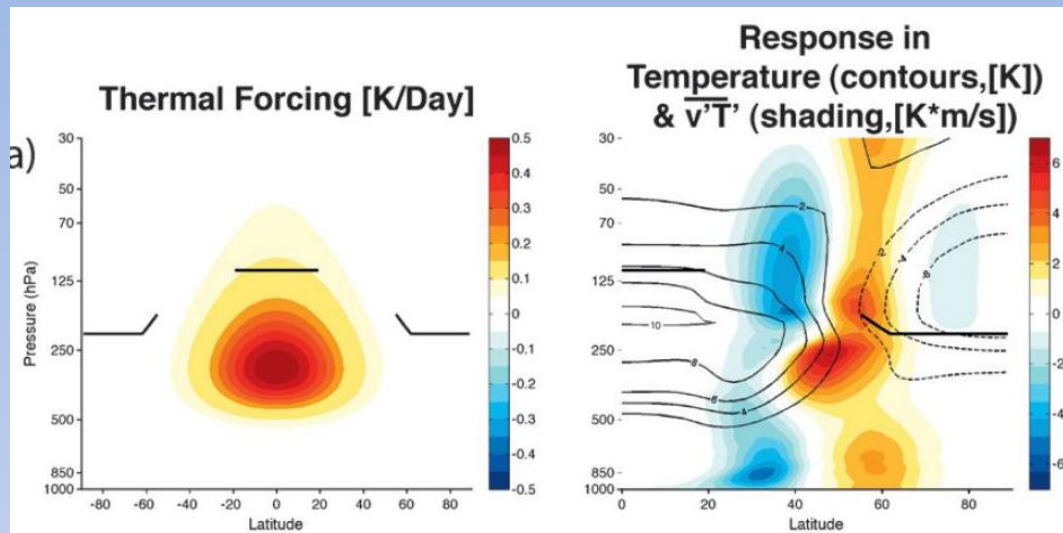
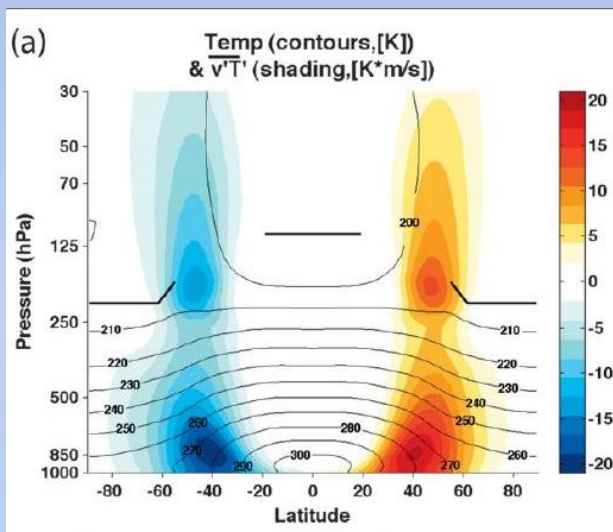
Equator-to-pole temperature differences

North Atlantic: 10W-60W



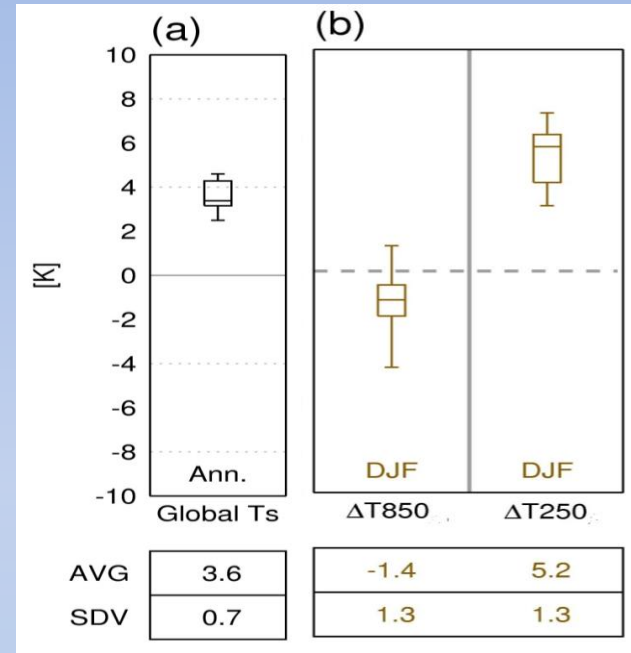
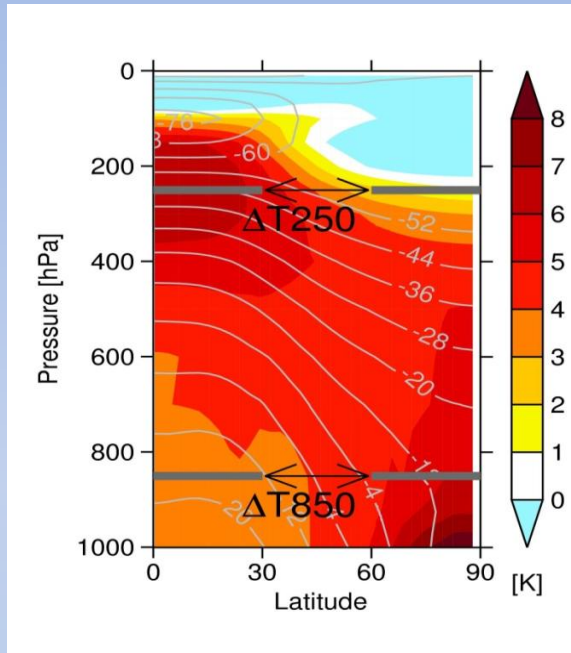
	Winter		
	Global Ts	Tropical Ts (ATL)	Polar Ts (ATL)
$\Delta T_{850_{ATL}}$	-0.13	0.10	-0.81
$\Delta T_{250_{ATL}}$	0.66	0.83	0.04

Idealised GCM study: Butler et al (2010)



Equator-to-pole temperature differences

North Atlantic: 10W-60W



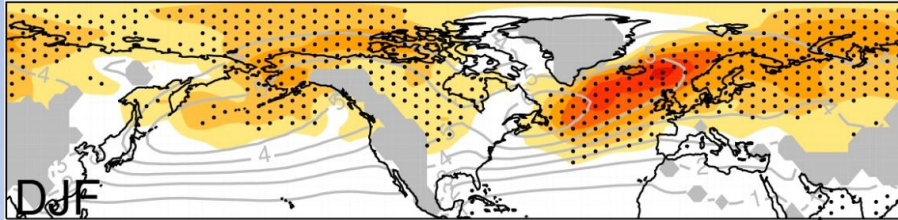
Perform a simple linear regression across the ensemble:

$$ST_{\text{resp},i} = \alpha + \beta \Delta T_{\text{resp},i} + \epsilon_i$$

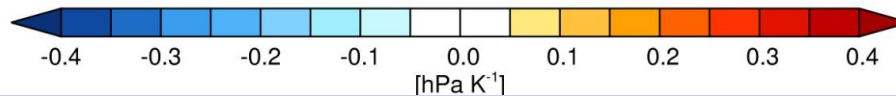
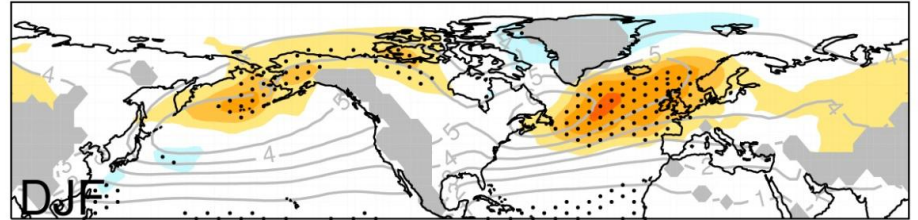
Results

Regression slopes (shading) and significance of correlation (stippling, $p=0.05$)

(a) winter $\Delta T_{850_{ATL}}$ regression slope



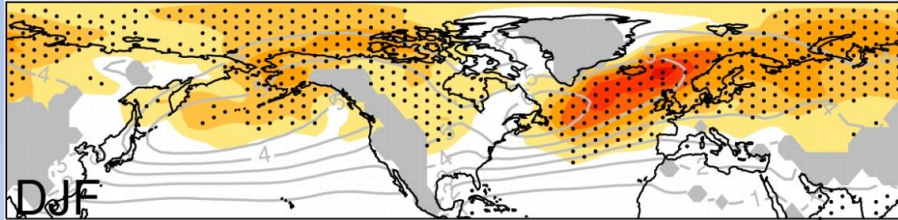
(b) winter $\Delta T_{250_{ATL}}$ regression slope



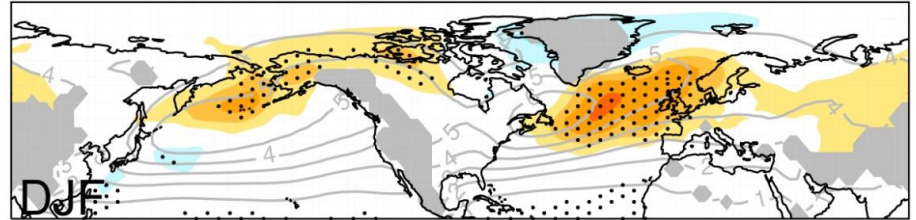
Results

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(a) winter $\Delta T_{850_{ATL}}$ regression slope

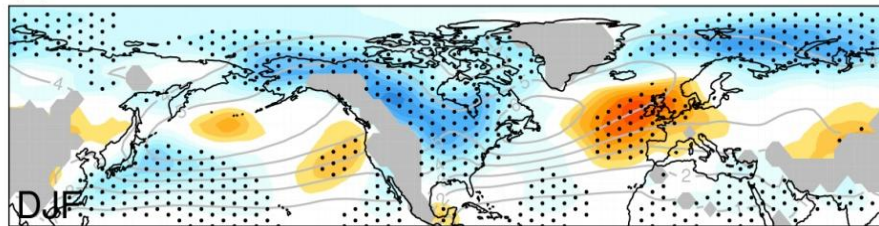


(b) winter $\Delta T_{250_{ATL}}$ regression slope

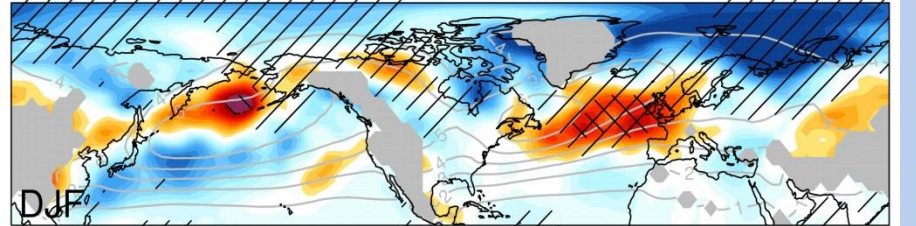


Comparison to mean storm track response:

(a) winter multi-model mean



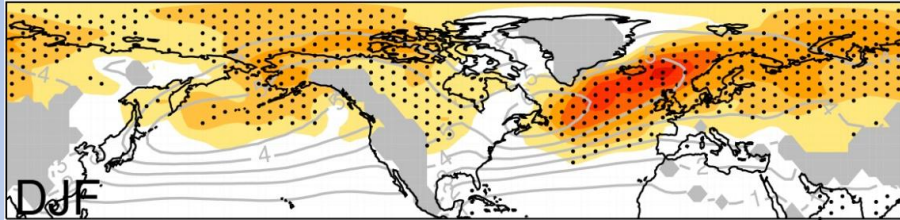
winter MLR (ATL):



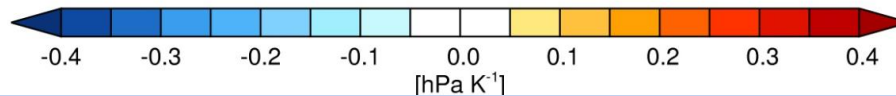
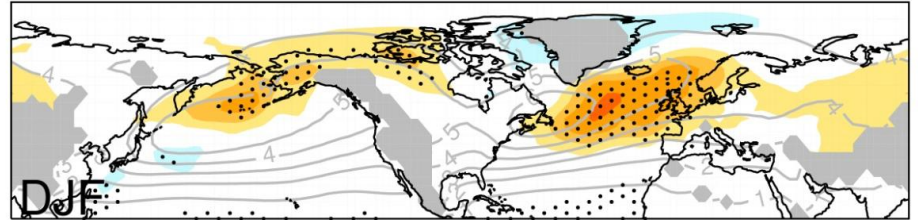
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Regression slopes (shading) and significance of correlation (stippling, $p=0.05$)

(a) winter $\Delta T_{850_{ATL}}$ regression slope

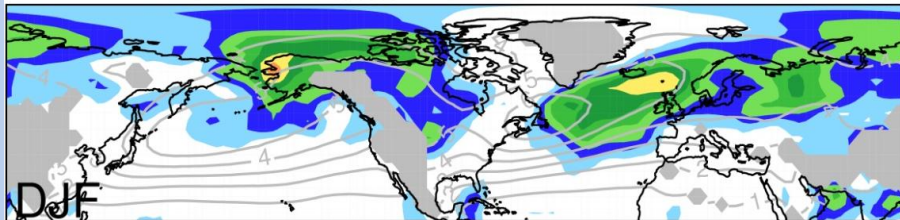


(b) winter $\Delta T_{250_{ATL}}$ regression slope

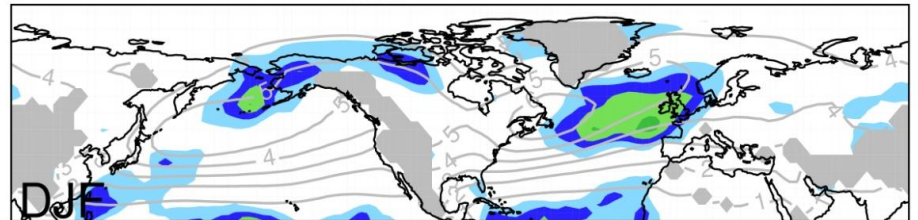


Fraction of inter-model variance 'explained':

(c) winter $\Delta T_{850_{ATL}}$ FVE



(d) winter $\Delta T_{250_{ATL}}$ FVE



Summary - North Atlantic

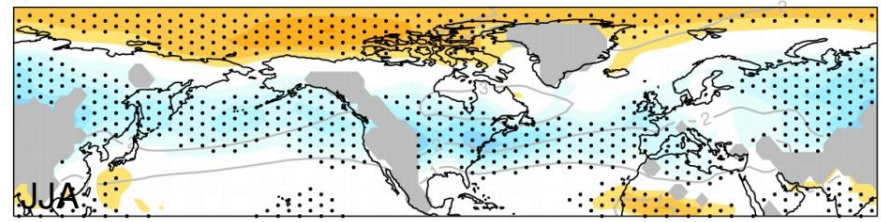
- There are **large regions with significant correlation** for both the ΔT_{850} and ΔT_{250} temperature differences.
- The **regression slopes** in these regions are mostly **positive**, suggesting the storm track responses are **driven by the responses of the baroclinicity**.
- The **impact of ΔT_{850}** on the multi-model mean storm track response is **negative** across most of the hemisphere, whereas the **impact of ΔT_{250}** is **positive** but confined to the ocean basins.
- Together, the two linear regression maps qualitatively **capture the spatial pattern of the multi-model mean response**.
- The FVE by ΔT_{850} is over **50%** in the North Atlantic and Norwegian Sea and by ΔT_{250} is over **30%** in the North Atlantic but small elsewhere.
- Suggests that there is potential to **reduce the spread** in the storm track responses by **constraining the relative strengths** of the warming in the tropics and polar regions.

Other regions

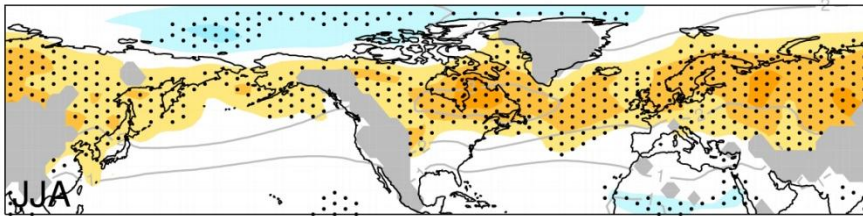
A similar analysis has been performed for both summer and winter of all the extratropical storm track regions.

e.g. NH summer

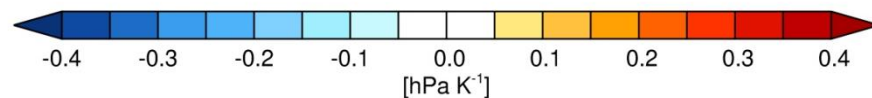
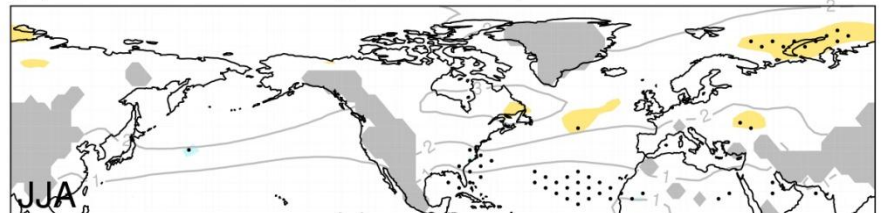
(c) summer multi-model mean



(c) summer ΔT_{850} regression slope



(d) summer ΔT_{250} regression slope

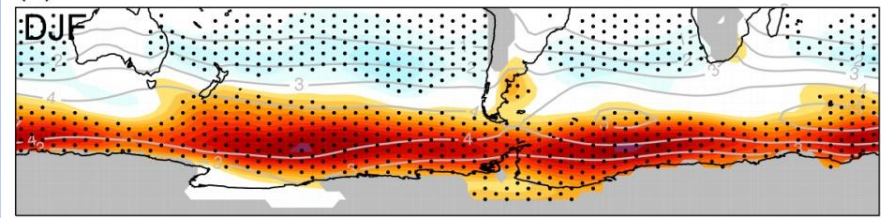


Other regions

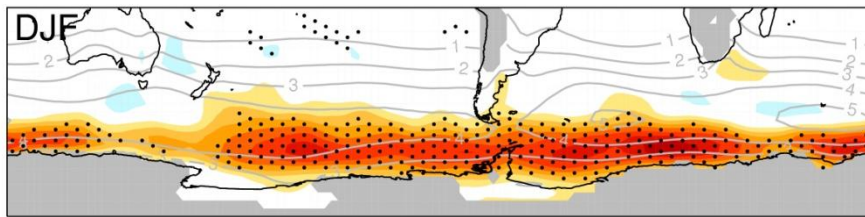
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e.g. SH summer

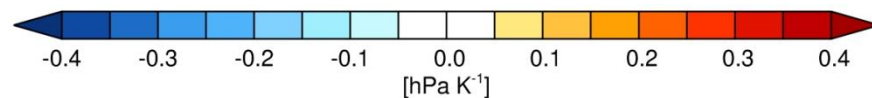
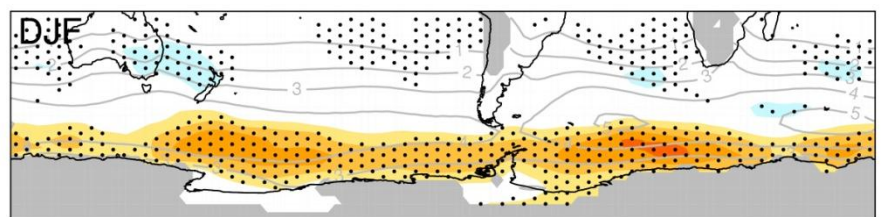
(c) summer multi-model mean



(c) summer ΔT_{850} regression slope



(d) summer ΔT_{250} regression slope



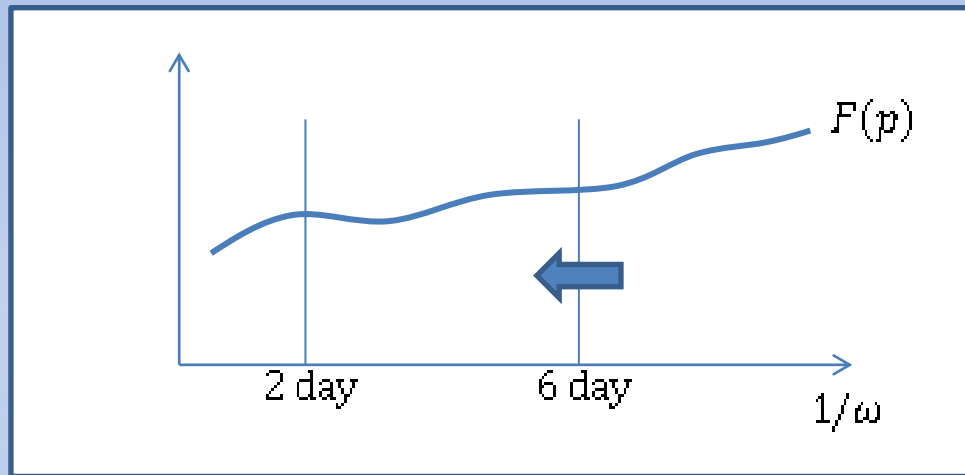
Summary - continued

- The **North Atlantic winter is unique** in that **both the ΔT_{850} and ΔT_{250} regressions are needed to capture the pattern of the mean response**. This more complex behaviour may go some way towards explaining the particularly large inter-model spread in the North Atlantic region.
- One limitation of this study is that the **causality of the correlations cannot be determined**. It is not clear whether the storm tracks respond directly to the **equator-to-pole temperature difference**, or instead to more **local baroclinicity changes** (e.g. SST, sea-ice or land-sea contrast changes) which may themselves be correlated with the equator-to-pole temperature difference.

Thank you for listening

What about Doppler effects?

A problem for bandpass filtered storm track measures

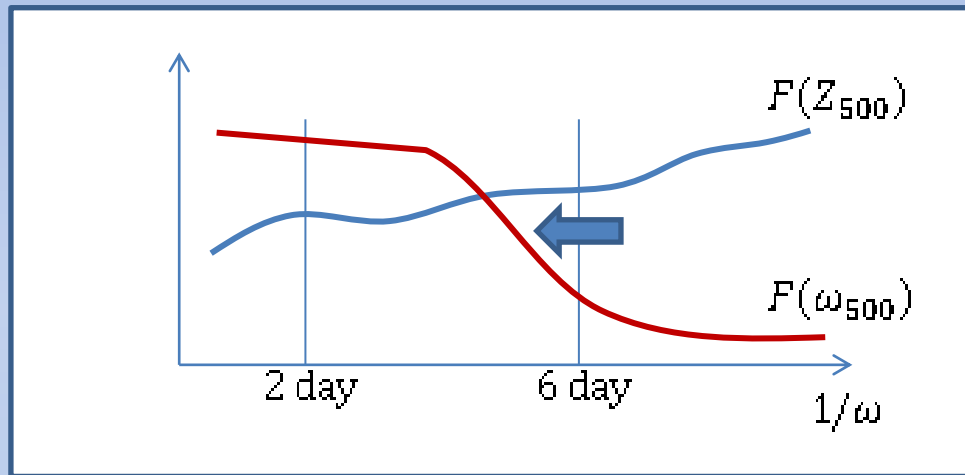


e.g. Burkhardt & James (2006)

- Investigate NAO/storm track relationship
- Use an 'extended EOF analysis' to correct for Doppler shifting

What about Doppler effects?

A problem for bandpass filtered storm track measures



e.g. Chang (2009)

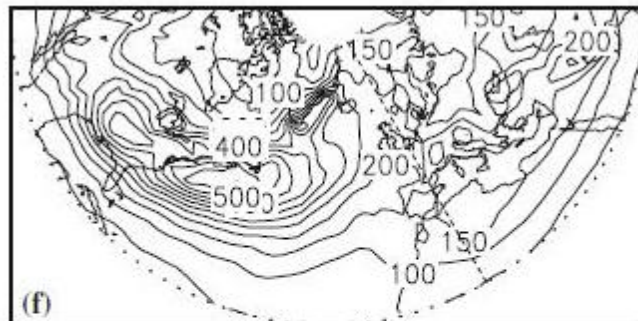
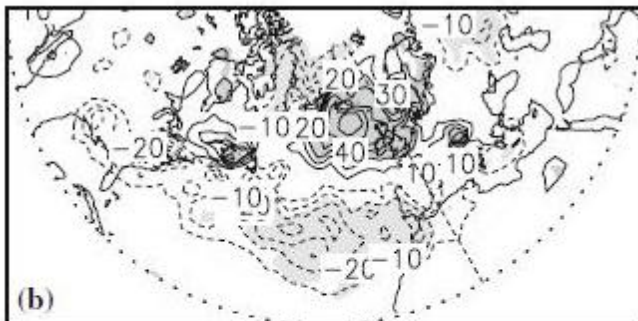
- Also investigate NAO/storm track relationship
- Use multiple storm track measures to understand problem

What about Doppler effects?

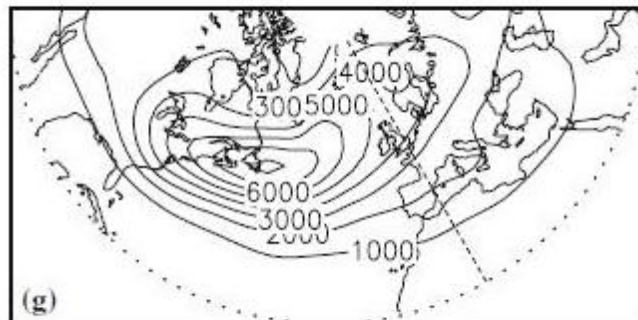
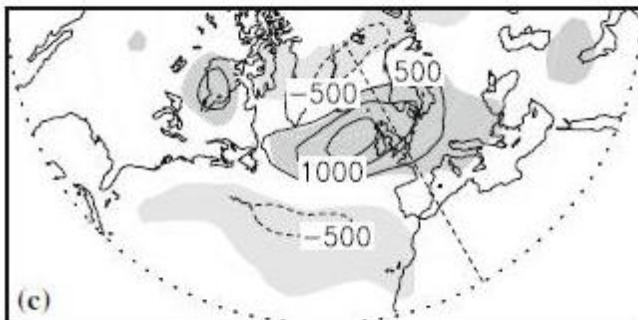
NAO regression

DJF mean

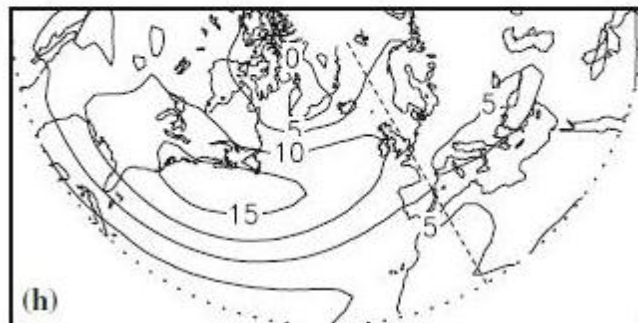
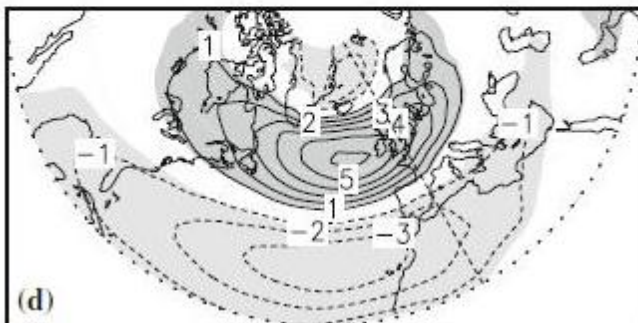
2 to 6 day ω_{500}



2 to 6 day Z_{500}



mean U_{700}

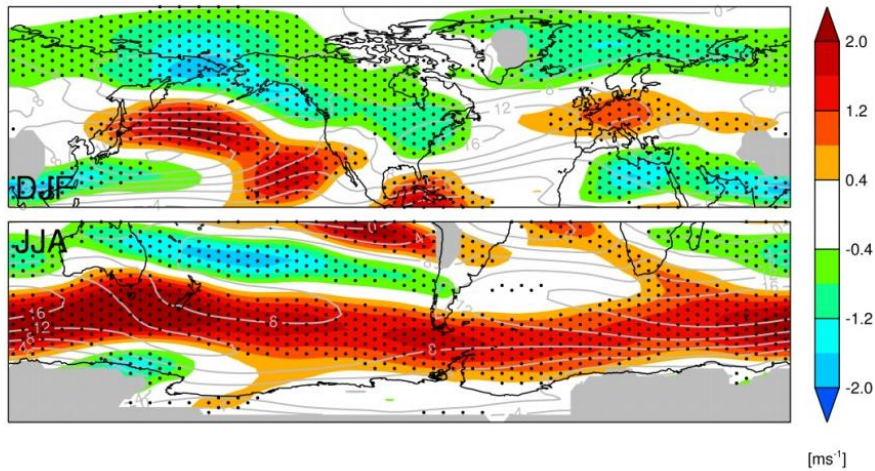


Chang (2009)

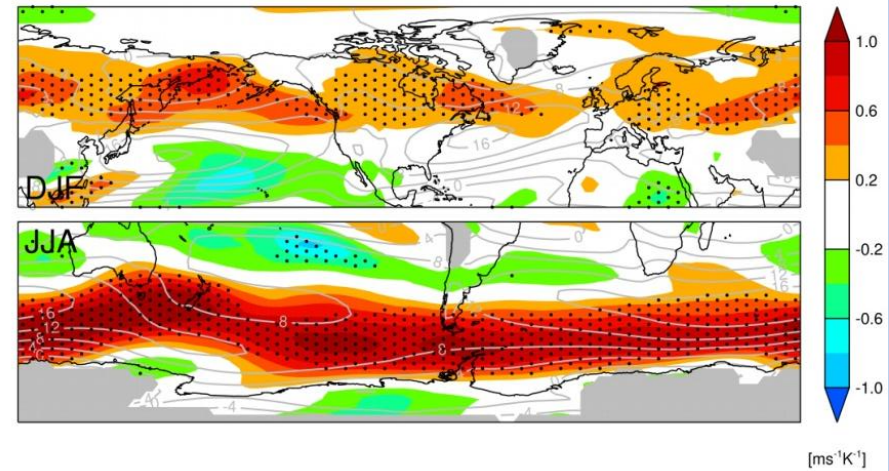
What about Doppler effects?

Could this be causing the storm track / temperature gradient relationship in the CMIP5 responses?

winter ua3 ensemble mean response



winter ua3 T850 gradient regression slope

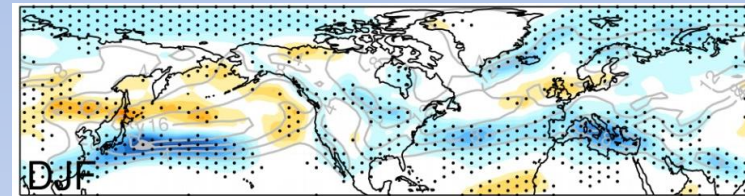
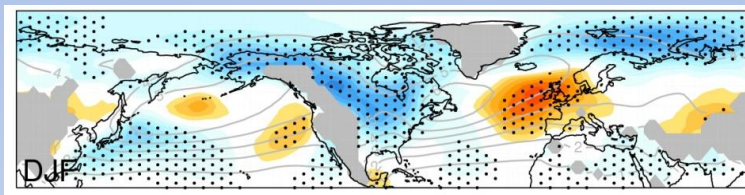


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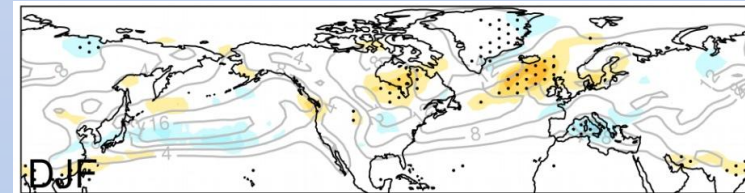
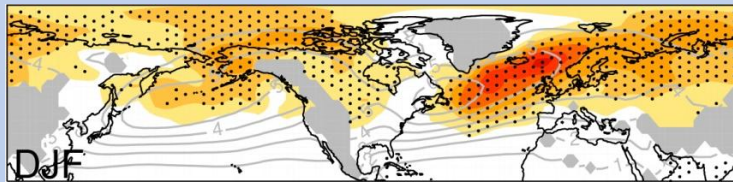
2-6 day MSLP standard deviation

Feature tracking using 850 hPa vorticity

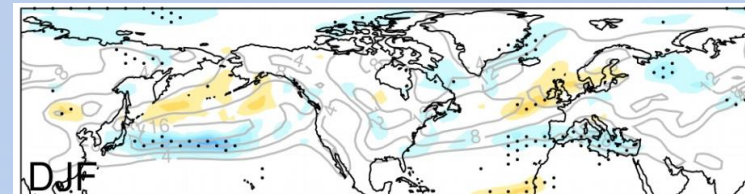
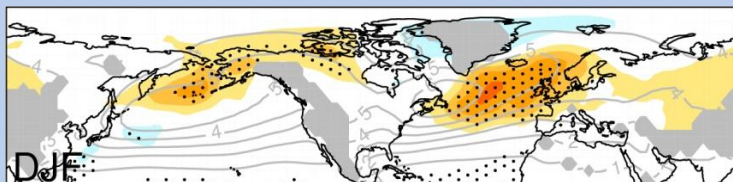
Ensemble mean response



T850 regression



T250 regression

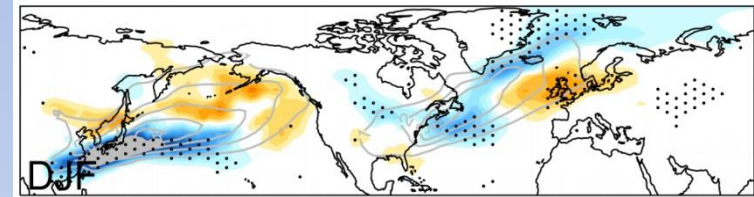
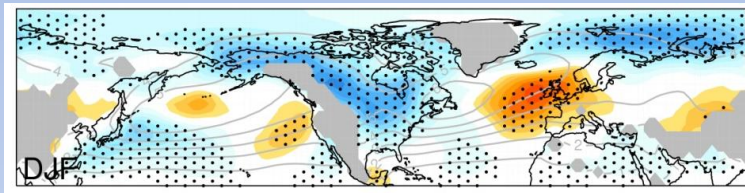


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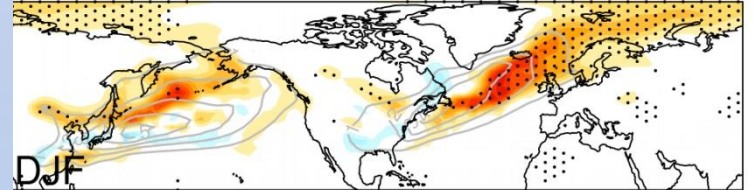
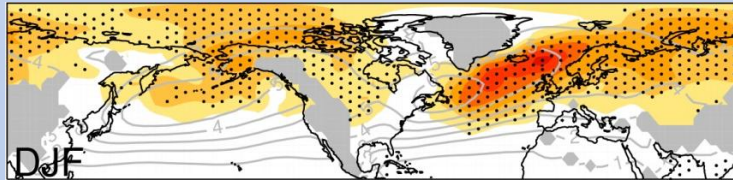
2-6 day MSLP standard deviation

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T850 regression



T250 regression

