COST Action Meeting 20-22 March 2012



Diagnosis of convective parameterisation schemes in extratropical cyclones

Oscar Martínez-Alvarado Bob Plant

Department of Meteorology University of Reading

DIAMET project



DIAbatic influences on Mesoscale structures in ExTratropical storms

- Consortium constituted by four UK universities (Manchester, Leeds, Reading and East Anglia) and the Met Office
- Three Work Packages
 - WP A. Structure of mesoscale anomalies and their wide-scale consequences
 - WP B. Physical processes and their parameterisation
 - WP C. Predictability

DIAMET project



DIAbatic influences on Mesoscale structures in ExTratropical storms

- Consortium constituted by four UK universities (Manchester, Leeds, Reading and East Anglia) and the Met Office
- Three Work Packages
 - WP A. Structure of mesoscale anomalies and their wide-scale consequences
 - WP B. Physical processes and their parameterisation
 - 1. Improving convective parameterisation
 - 2. Air-sea fluxes and their influence on storm development
 - 3. Microphysical processes
 - WP C. Predictability

- Tracers (affected by advection) for potential temperature (θ), specific humidity
 (q), cloud liquid water (q_{cl}) and cloud ice content (q_{cf})
- The variables of interest (θ , q, q_{cl} and q_{cf}) are decomposed in the following way:

$$\varphi = \varphi_0 + \Delta \varphi \tag{1}$$

• In turn the total increment to the variable is decomposed as follows:

$$\Delta \varphi(x,t) = \sum_{i=\text{proc}} \Delta \varphi_i(x,t)$$
(2)

proc = {parameterised processes}

- These processes depend on the variable under consideration. For example,
 - For potential temperature:
 - Microphysics, mixing in BL, latent heating in BL, convection, radiation
 - For specific humidity:
 - Microphysics, boundary layer, convection





• For each tracer there is an evolution equation of the form

$$\frac{D\varphi_0}{Dt} = \frac{\partial \varphi_0}{\partial t} + \mathbf{v} \cdot \nabla \varphi_0 = 0 \tag{3}$$

• The initial field (ϕ_0) satisfies the following equation

$$\frac{D\Delta\varphi_i}{Dt} = \frac{\partial \Delta\varphi_i}{\partial t} + \mathbf{v} \cdot \nabla \Delta\varphi_i = S_{\varphi_i}$$
⁽⁴⁾

• Rewriting equation (3) and (4) as

$$\frac{\partial \varphi_0}{\partial t} = -\mathbf{v} \cdot \nabla \varphi_0 \tag{5}$$

$$\frac{\partial \Delta \varphi_i}{\partial t} = -\mathbf{v} \cdot \nabla \Delta \varphi_i + S_{\varphi_i} \tag{6}$$



• For each tracer there is an evolution equation of the form

$$\frac{D\varphi_0}{Dt} = \frac{\partial \varphi_0}{\partial t} + \mathbf{v} \cdot \nabla \varphi_0 = 0$$
(3)

- The initial field (ϕ_0) satisfies the following equation

$$\frac{D\Delta\varphi_{i}}{Dt} = \frac{\partial \Delta\varphi_{i}}{\partial t} + \mathbf{v}\cdot\nabla\Delta\varphi_{i} = S_{\varphi_{i}}$$
(4)
• Rewriting equation (3) and (4) as
$$\frac{\partial \varphi_{0}}{\partial t} = -\mathbf{v}\cdot\nabla\varphi_{0}$$
(5)
$$\frac{\partial \Delta\varphi_{i}}{\partial t} = -\mathbf{v}\cdot\nabla\Delta\varphi(+S_{\varphi_{i}})$$
Affected by
sources/sinks of
the relevant
variable
(6)



• The evolution equation for the relevant variables can then be written as

$$\frac{\partial \varphi}{\partial t} = \frac{\partial \varphi_0}{\partial t} + \frac{\partial \Delta \varphi_i}{\partial t} = -\mathbf{v} \cdot \nabla \varphi_0 - \mathbf{v} \cdot \nabla \sum_{i = \text{proc}} \Delta \varphi_i + \sum_{i = \text{proc}} S_{\varphi_i}$$
(7)



• The evolution equation for the relevant variables can then be written as





• The evolution equation for the relevant variables can then be written as





Data and model

- Case from DIAMET first field campaign:
 - 30 September 2011
 - Low-pressure system centred to the south-west of Iceland
 - Long trailing active cold front
- Model:
 - Met Office Unified Model (MetUM) version 7.3
 - North-Atlantic—Europe (NAE) domain
 - Grid spacing 0.11° (~12 km)
 - 38 vertical levels (lid ~40 km)
 - (MetUM Modified) Gregory—Rowntree convection scheme

DIAMET field campaign 0600 UTC 30 September 2011





850-hPa equivalent potential temperature

30 September 2011 0600 UTC

Mean sea-level pressure

DIAMET field campaign 0600 UTC 30 September 2011



Model-derived OLR

30 September 2011 0600 UTC



Model derived Total precipitation



850-hPa equivalent potential temperature

0600 UTC 30 September 2011 Comparison with observations



Radar image



Model derived Total precipitation



0600 UTC 30 September 2011 Convective – Large-scale rain split





Total precipitation

Convective rain

0600 UTC 30 September 2011 Convective – Large-scale rain split







Preliminary results

- Integration starting on 1200 UTC 28 September 2011
- The following slides show cumulative increments on 0600 UTC 30
 September 2011 (T+42)

Sources/sinks of potential temperature 850 hPa (~1300 m height)









30 September 2011 0600 UTC, pressure - 850 hPa

30 September 2011 0600 UTC, pressure - 850 hPa



30 September 2011 0600 UTC, pressure = 850 hPa

Sources/sinks of potential temperature 850 hPa (~1300 m height)





Sources/sinks of potential temperature University of **Reading** 850 hPa (~1300 m height) 30 September 2011 Competing processes along the cold front and **Microphysics** Total change tion below the cloud head 850 bPa 30 September 2011 0600 UTC, pressure = 850 hP **BL** latent heat Convection BL mixing



Sources/sinks of potential temperature 600 hPa (~4000 m height) UTC. pressure = 600 hPa









30 September 2011 0600 UTC, pressure - 600 hPa



Sources/sinks of potential temperature 600 hPa (~4000 m height)





Sources/sinks of potential temperature 600 hPa (~4000 m height)





Sources/sinks of potential temperature 600 hPa (~4000 m height)



tion





Cloud head formed by air from the boundary layer

30 September 2011 0600 UTC, pressure - 600 hPa









Sources/sinks of potential temperature 250 hPa (~10000 m height)







30 September 2011 0600 UTC, pressure - 250 hPa



30 September 2011 0600 UTC, pressure - 250 hPa

30 September 2011 0600 UTC, pressure - 250 hPa

Convection BL latent heat

Sources/sinks of potential temperature 250 hPa (~10000 m height)









30 September 2011 0600 UTC, pressure - 250 hPa





30 September 2011 0600 UTC, pressure = 250 hPa



Sources/sinks of potential temperature





Concluding remarks



 One tools for the analysis of convection in large-scale atmospheric models have been presented:

Budget of energy (heating/cooling) and moisture (drying/moistening)

- It represents an integral view of the processes acting during a simulation (cumulative action/cumulative increments)
 - It is useful not only for the evaluation of convection but for a complete assessment of processes in the model
 - It does not provide instantaneous information about sources and sinks
- It can be coupled with other budget diagnostics such as that for potential vorticity

Concluding remarks



- Instantaneous information about the generation/destruction of the different variables can be obtained
 - By reformulating the approach to output instantaneous increments (perhaps averaged over a predetermined period) or
 - By combining the present approach with trajectories to study the changes due to parameterised processes along trajectories

Concluding remarks



- This method could be useful for
 - The analysis of balance between convective and large-scale achieved by
 - Different spatial model resolutions
 - Different parameterisation schemes (with a focus on convection schemes).
 - The analysis of a broader range of phenomena not restricted to convection parameterisation schemes