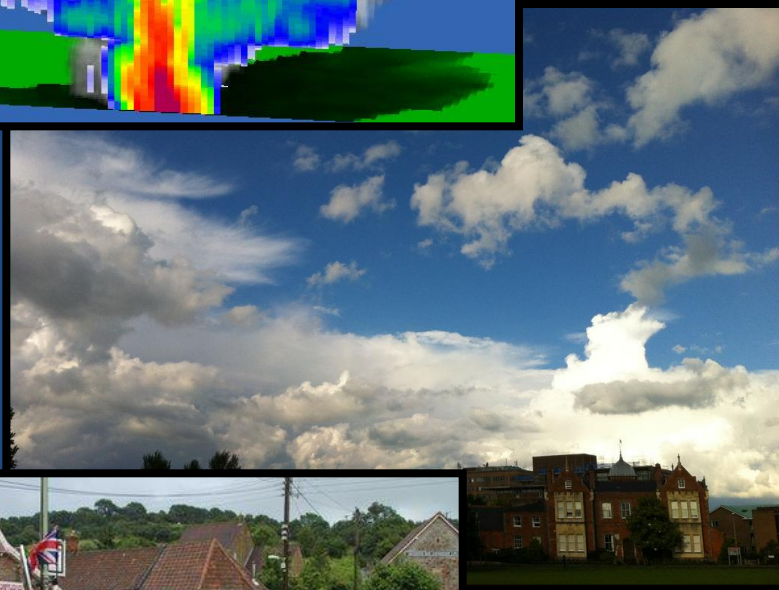
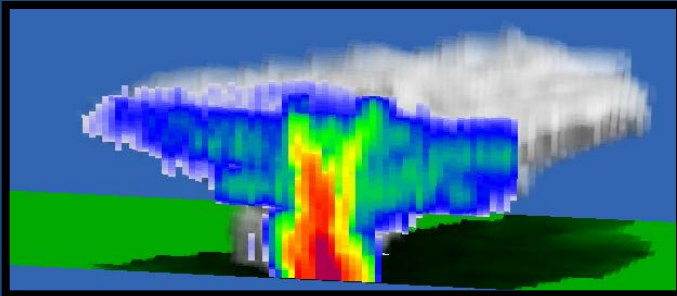


The Dynamical and Microphysical Evolution of Convective Storms

DYMECS

A statistical evaluation of convective storms in high-resolution Unified Model simulations



John Nicol

Robin Hogan
Thorwald Stein
Robert Plant
Peter Clark



University of
Reading

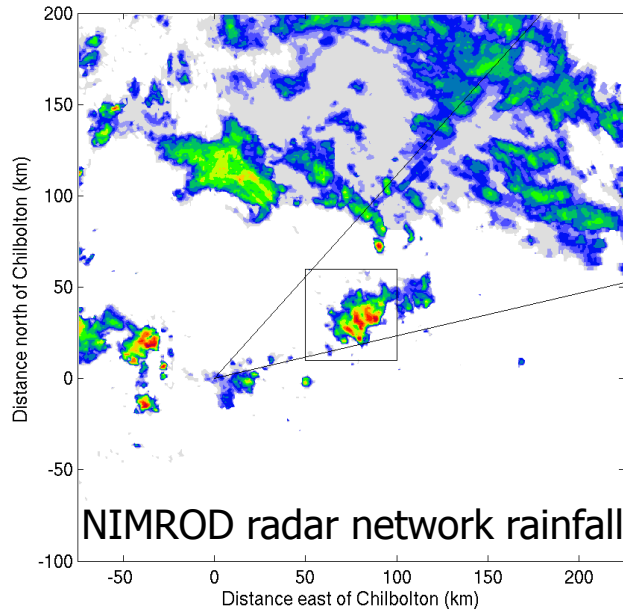
MetOffice@Reading:
Carol Halliwell
Kirsty Hanley
Humphrey Lean



Met Office

Chilbolton:
Mal Clarke
Alan Doo
Darcy Ladd

The DYMECS approach: beyond case studies

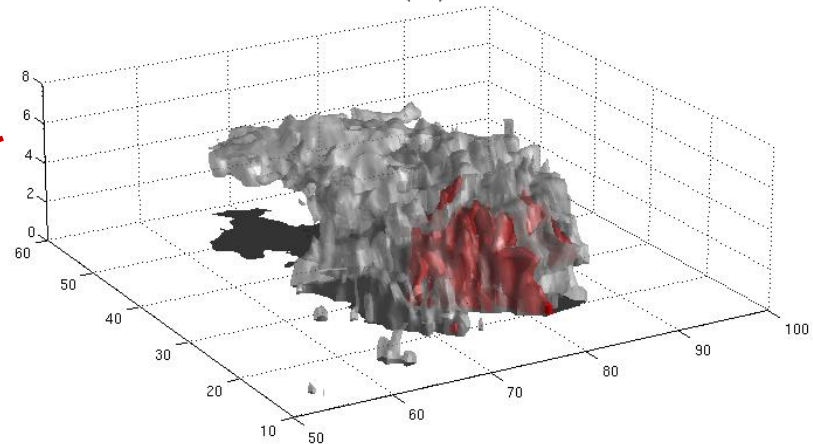
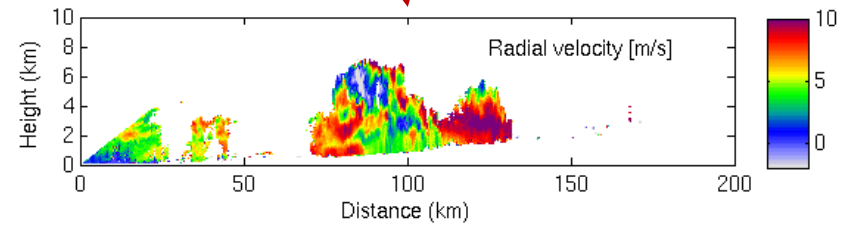
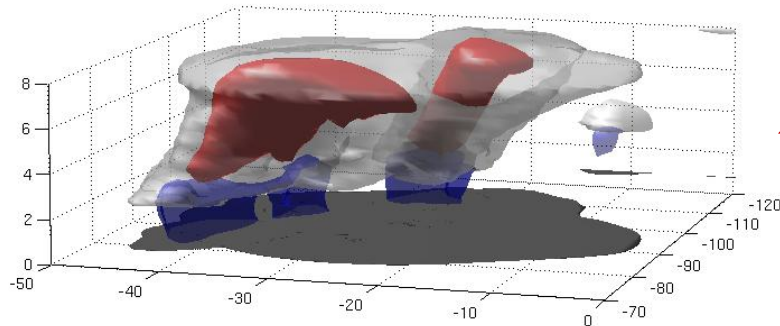


Track storms in real time
and automatically scan
Chilbolton radar



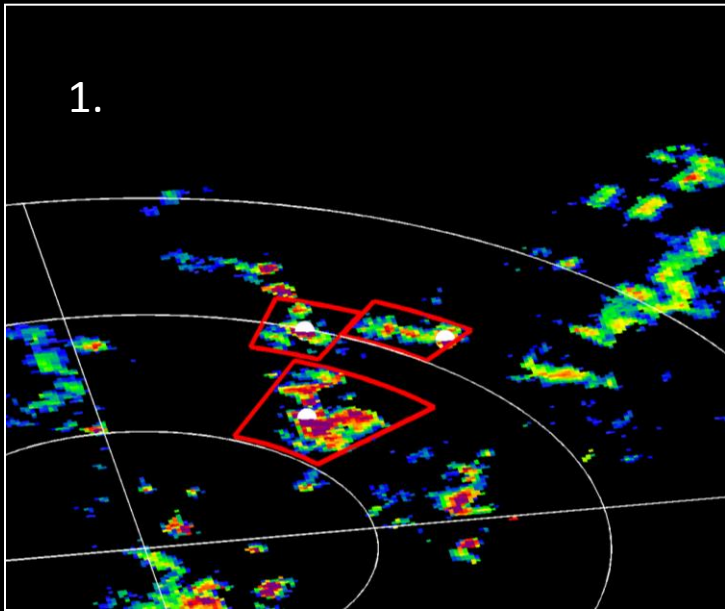
Derive properties of
hundreds of storms
on **~40 days**:
Vertical velocity
3D structure
Rain & hail
Ice water content
TKE & dissipation rate

Evaluate these properties in model varying:
Resolution
Microphysics scheme
Sub-grid turbulence parametrization



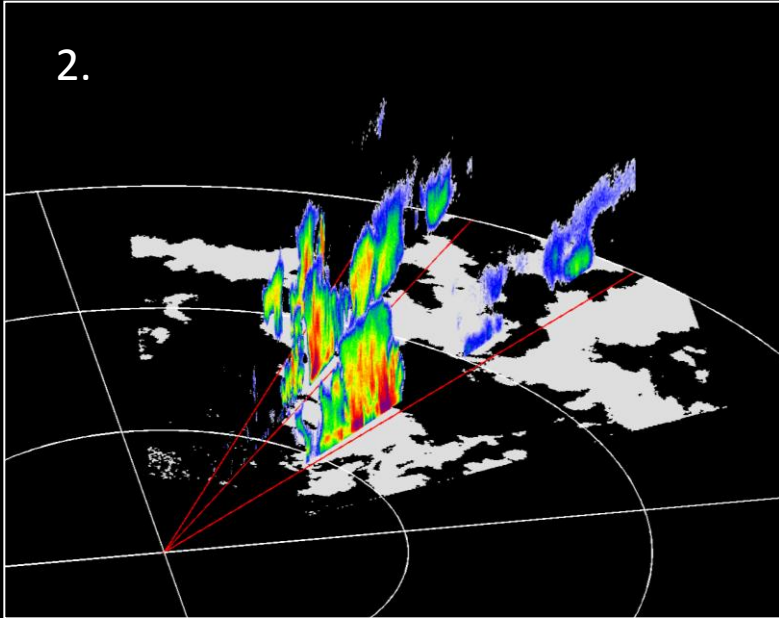
Automated scanning with CAMRa

1.

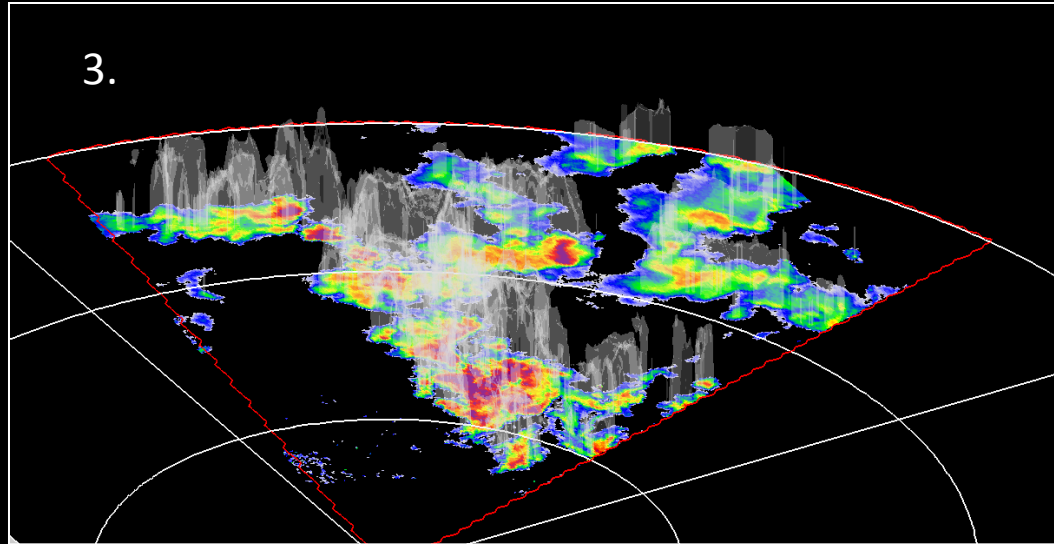


0. Storms were tracked in MetOffice rainfall radar data.
1. The scan scheduler prioritized storms by area and mean rainfall, with a preference for previously scanned storms and locations.
2. Sets of 4 RHIs were performed through locations of maximum rainfall.
3. A volume scan of 6-12 PPIs was performed through prioritized storms (either separately, or in a single volume if grouped closely in azimuth).

2.



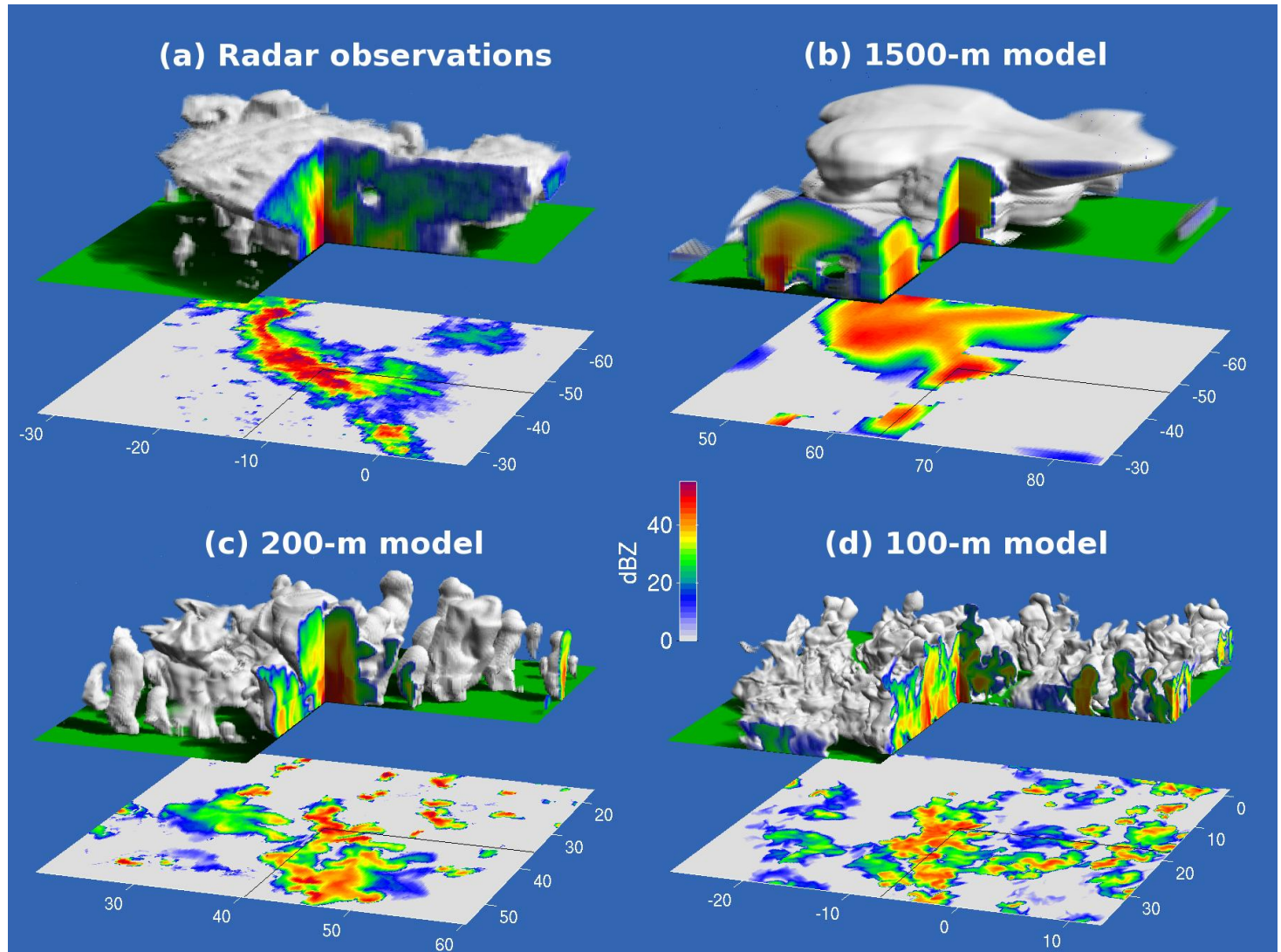
3.



3D visualisation of data

25 Aug 2012

Cutaway:
reflectivity
Surface:
rainrate
Shading:
extent of cloud



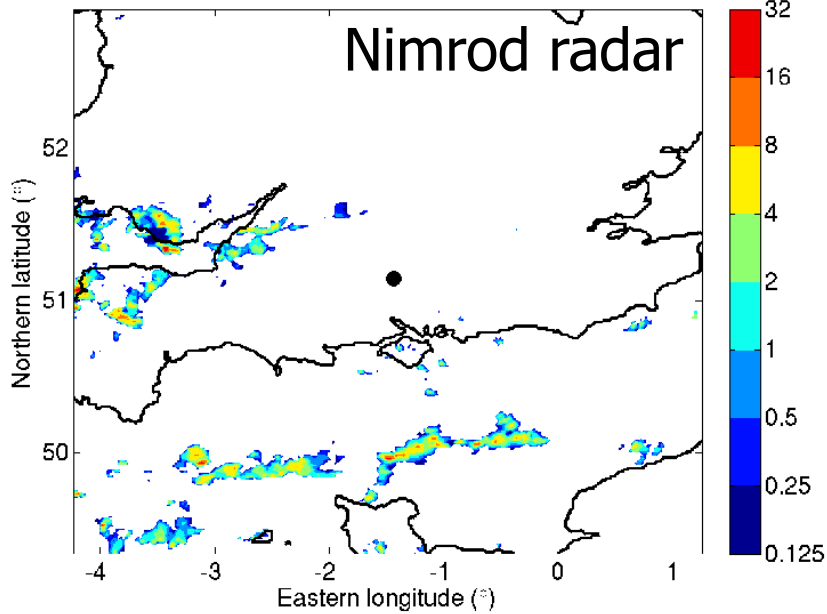
Evaluate convective storm characteristics in the Unified Model

Precipitation patterns and life cycles

Storm and updraught structure

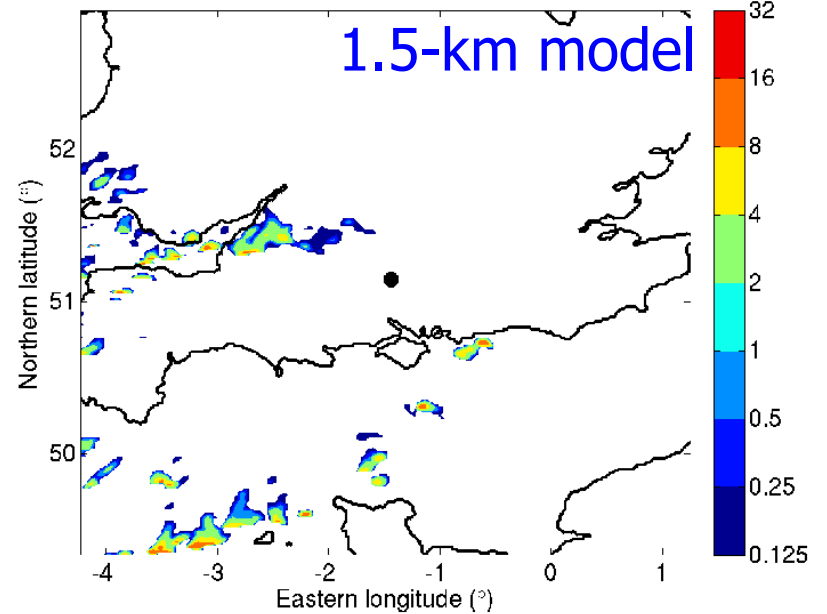
Rainfall rate for 7 : 10 UTC 07/08/11 - Nimrod

Nimrod radar



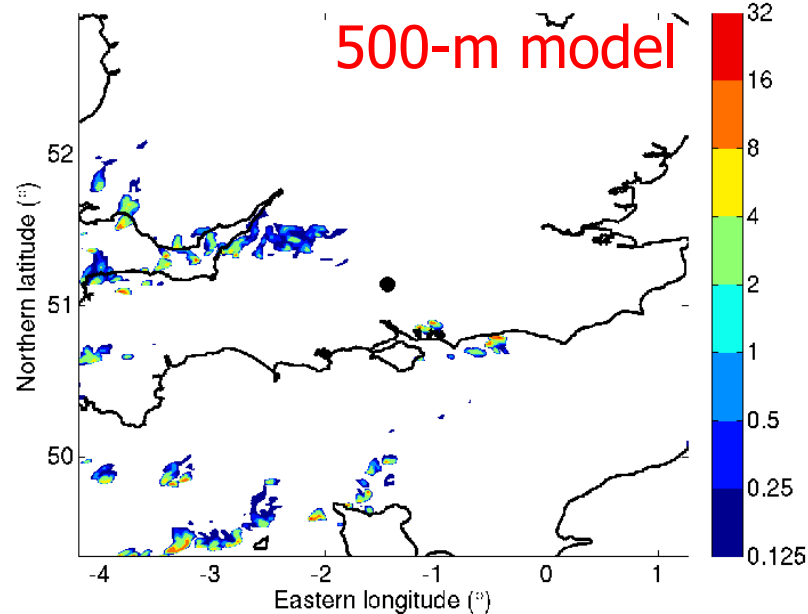
Rainfall rate for 7 : 10 UTC 07/08/11 - UKV

1.5-km model



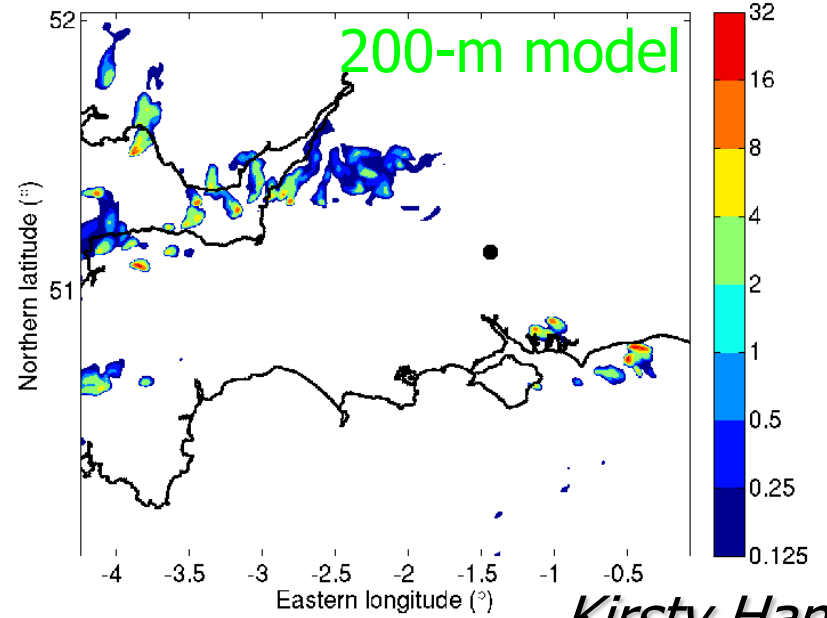
Rainfall rate for 7 : 10 UTC 07/08/11 - 500m smoothed

500-m model

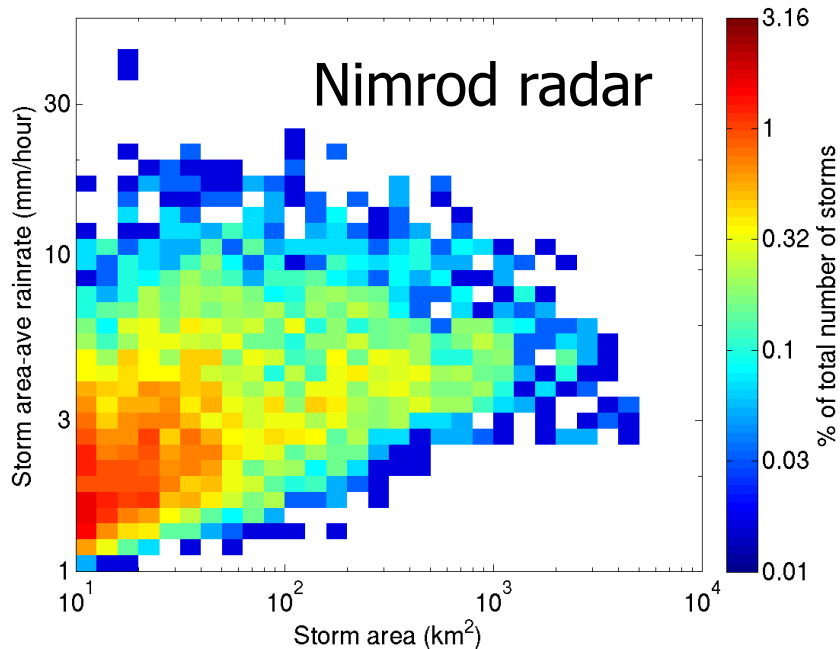


Rainfall rate for 7 : 10 UTC 07/08/11 - 200m smoothed

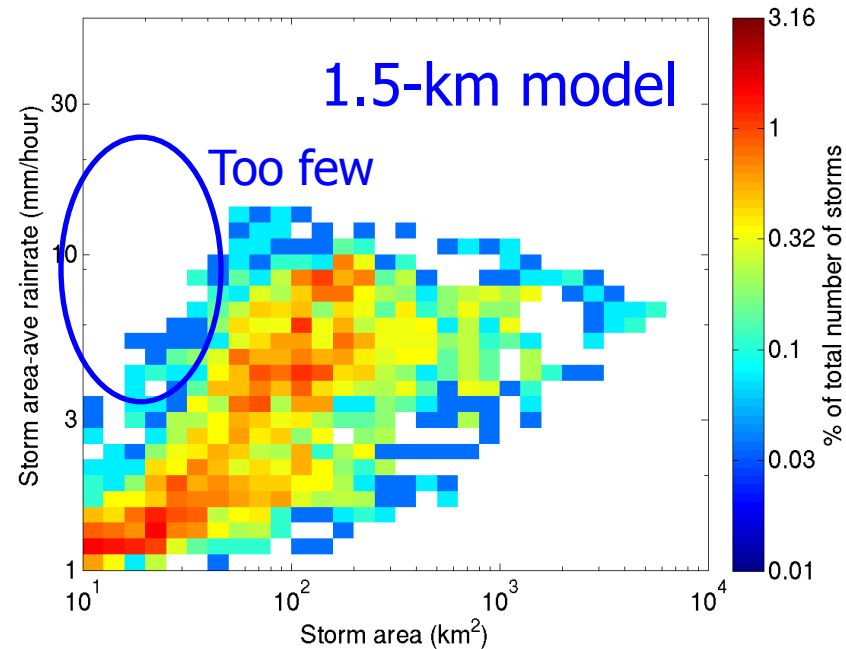
200-m model



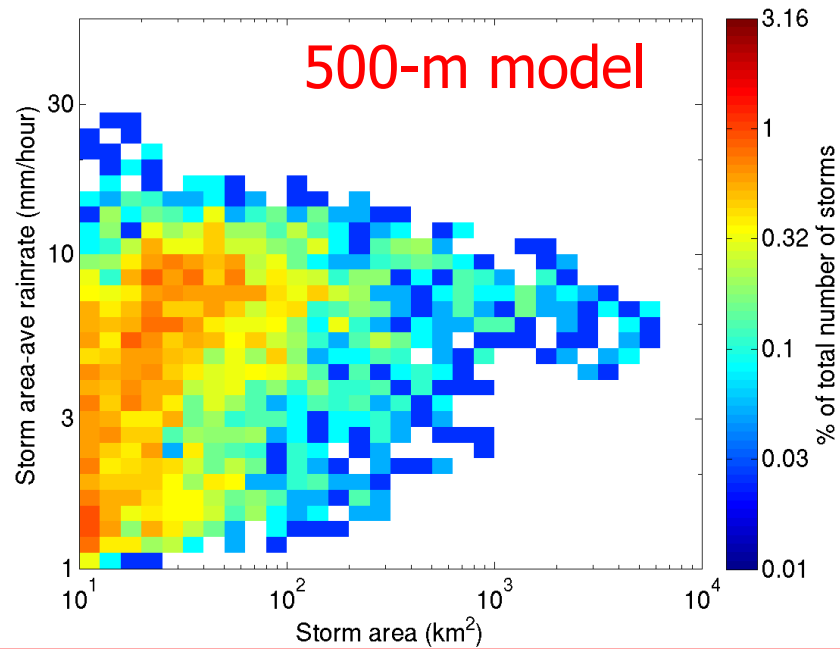
20110807 9 - 19 UTC 300 x 225 km R = 1 A= 10 - Nimrod



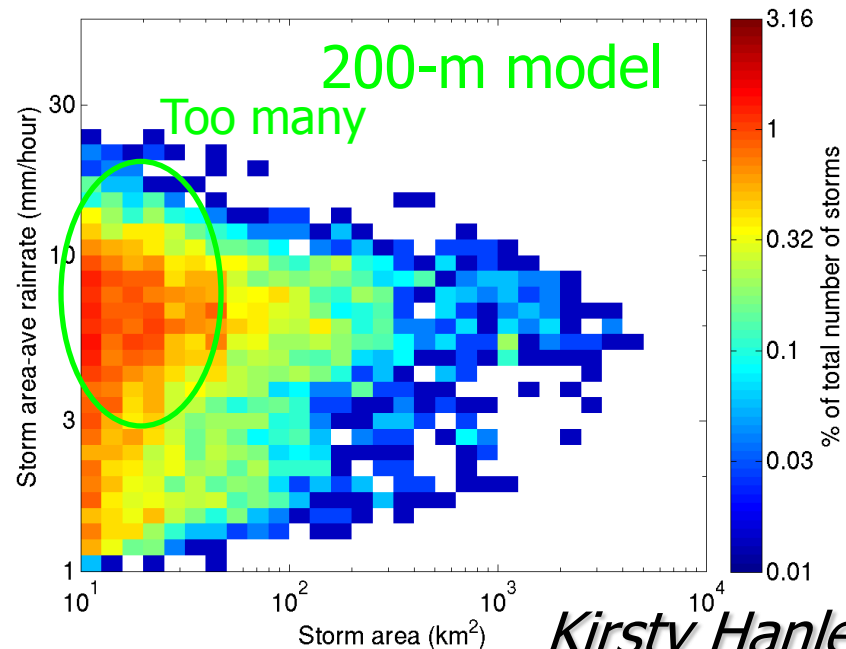
20110807 9 - 19 UTC 300 x 225 km R = 1 A= 10 - 1500m



20110807 9 - 19 UTC 300 x 225 km R = 1 A= 10 - 500m

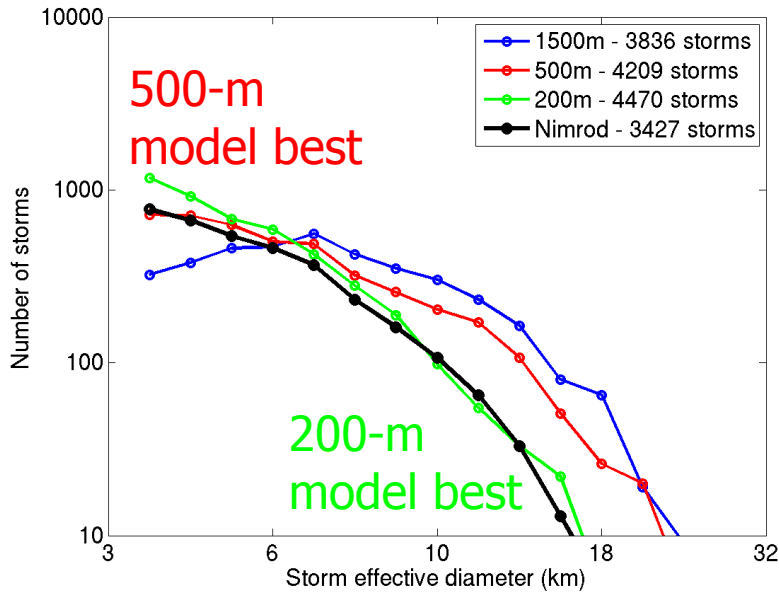
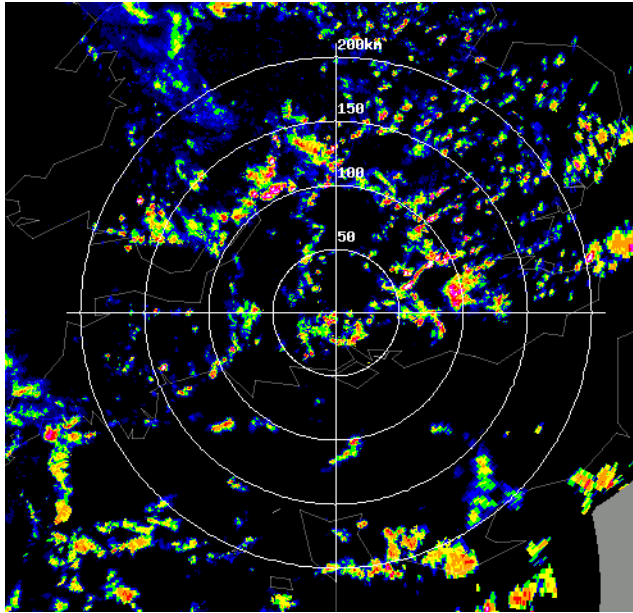


20110807 9 - 19 UTC 300 x 225 km R = 1 A= 10 - 200m

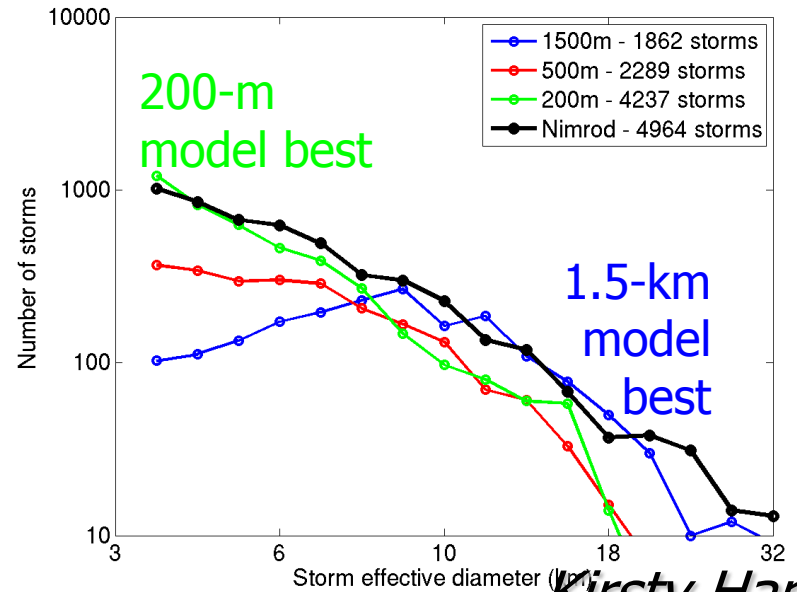
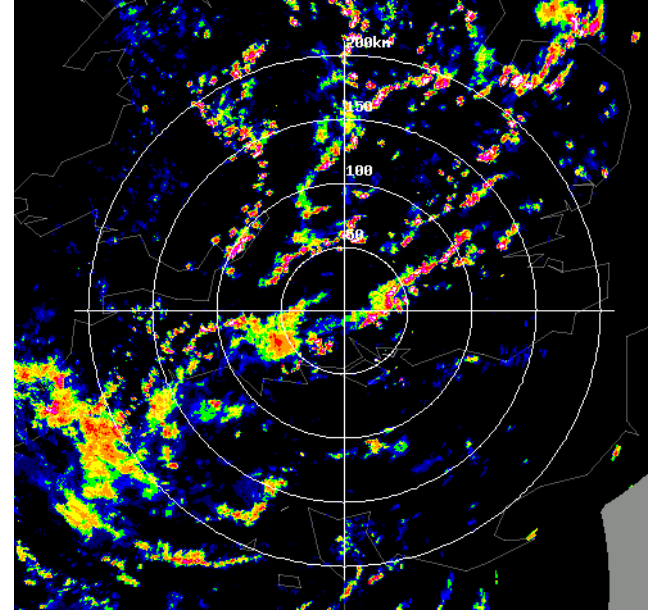


Kirsty Hanley

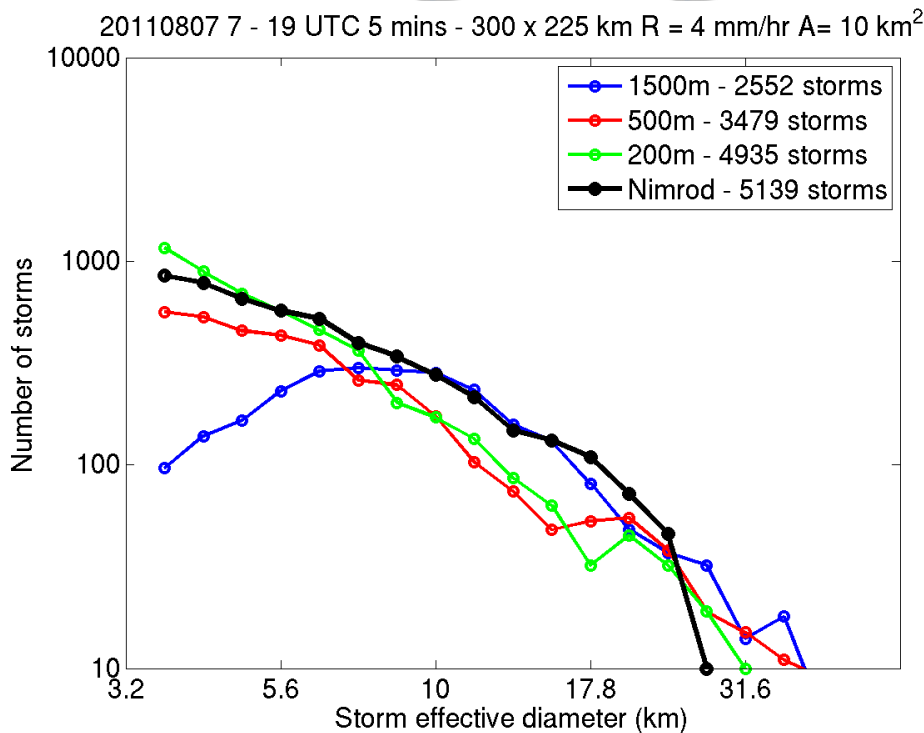
20 April 2012



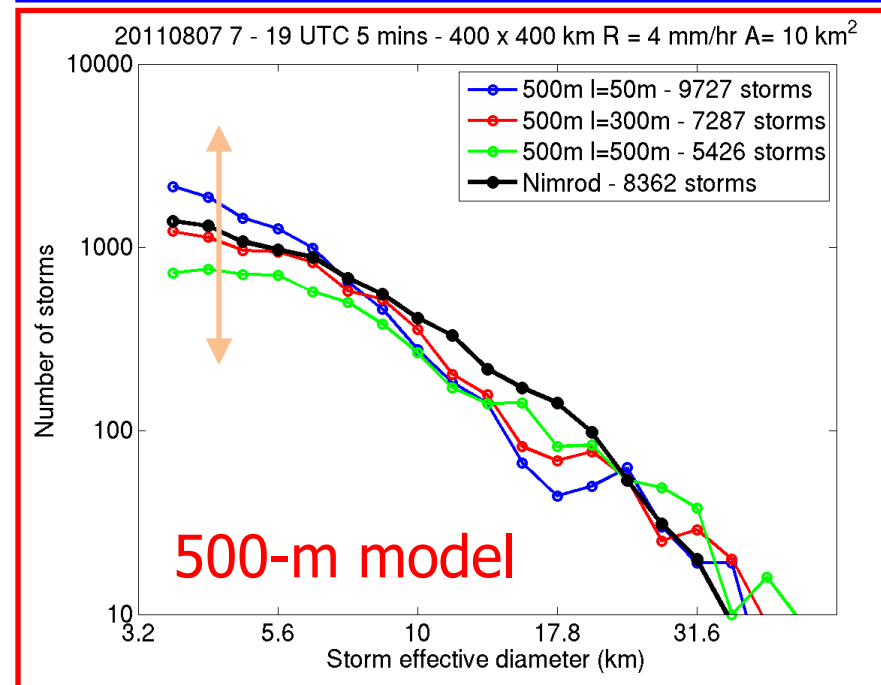
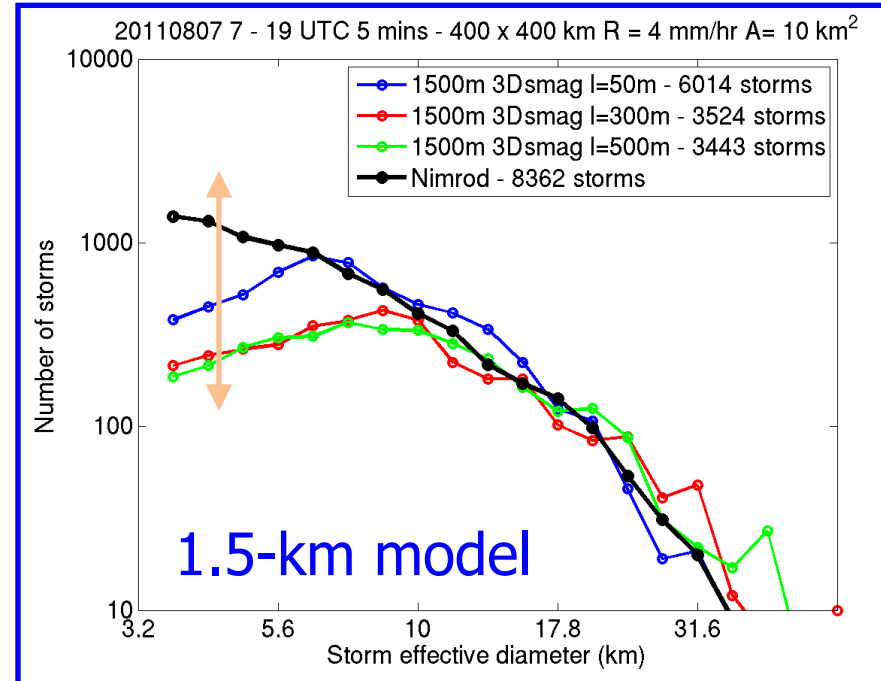
25 Aug 2012



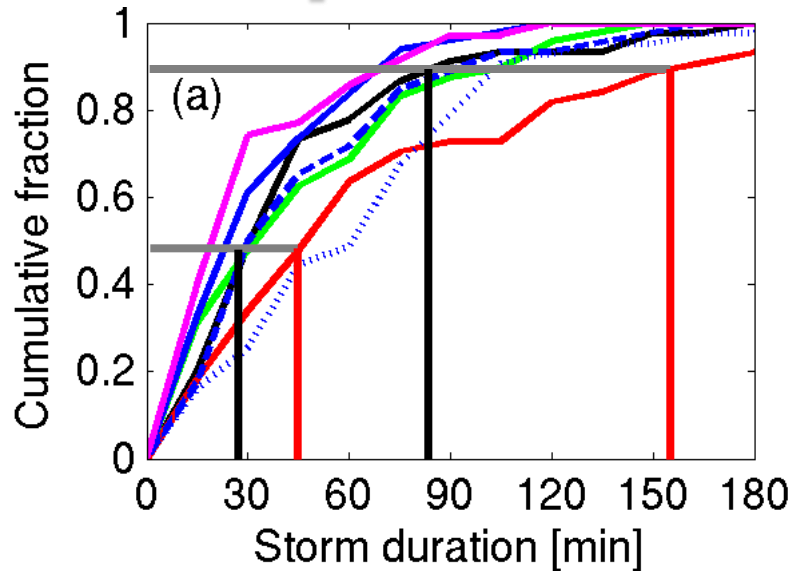
Storm size- Mixing length



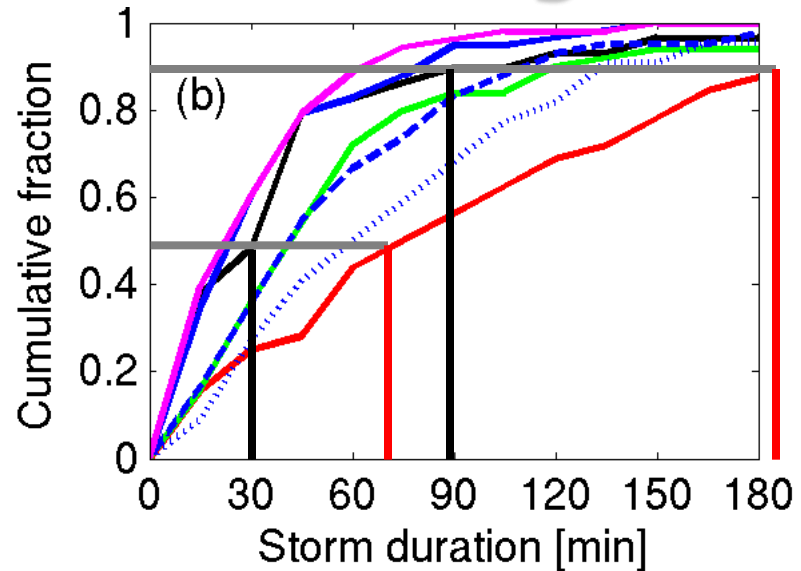
- Smagorinsky mixing length plays a key role in determining number of small storms



20 April 2012



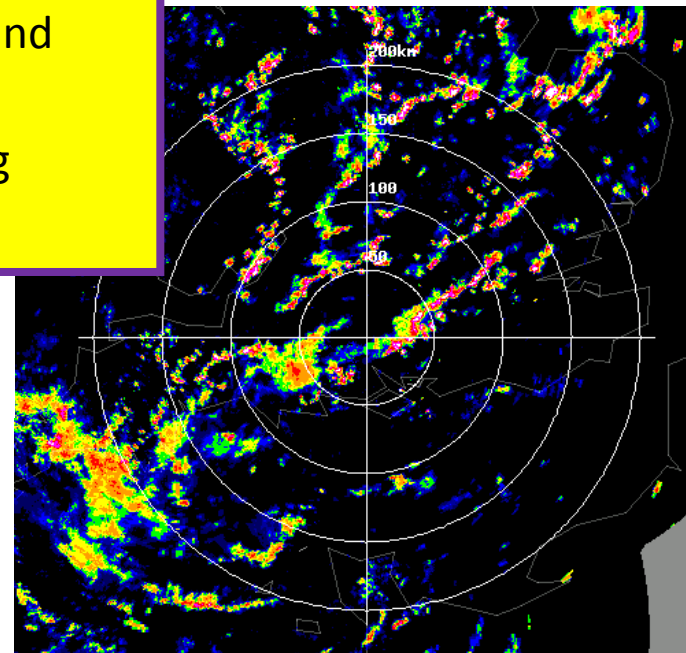
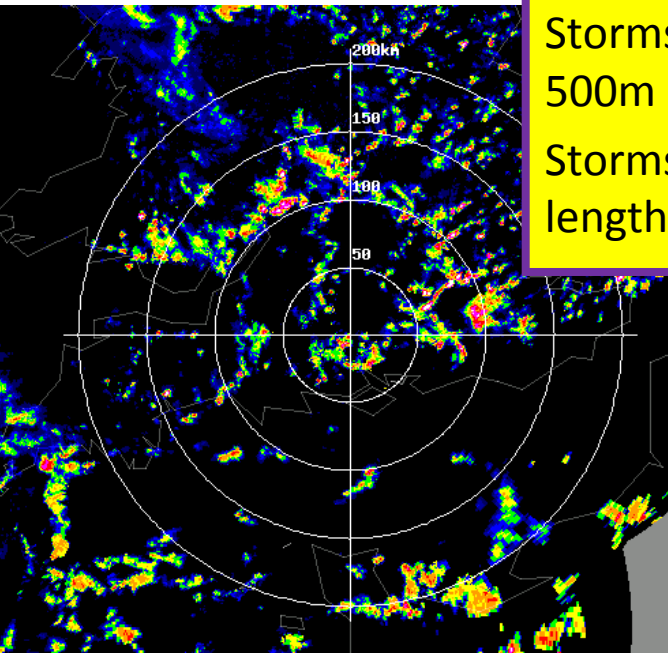
25 Aug 2012



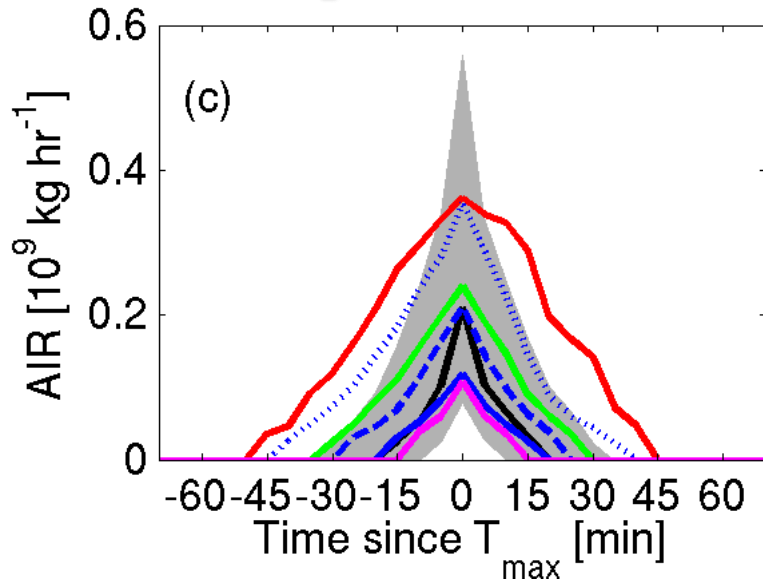
Storms last too long in the 1500m and 500m models.

Storms also last longer when mixing length is increased.

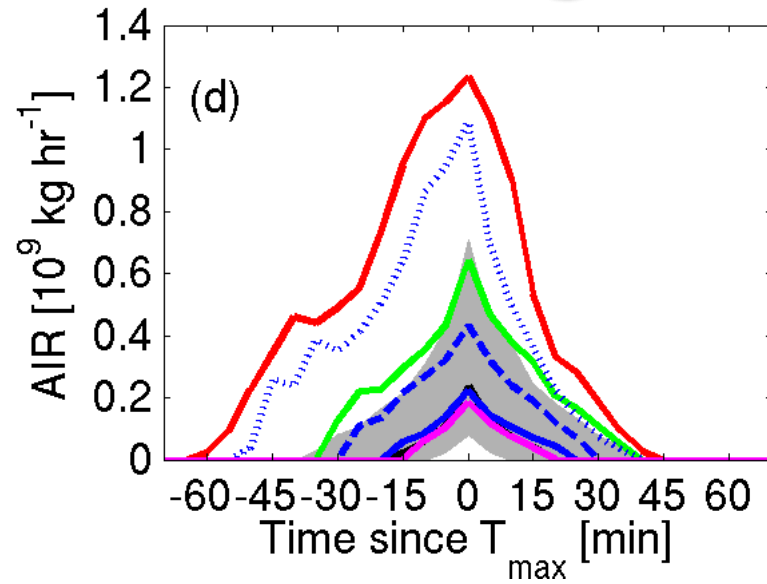
- 200m model, 40m mixing length
- 100m model
- - - 200m model, 100m mixing length
- ... 200m model, 300m mixing length



20 April 2012



25 Aug 2012

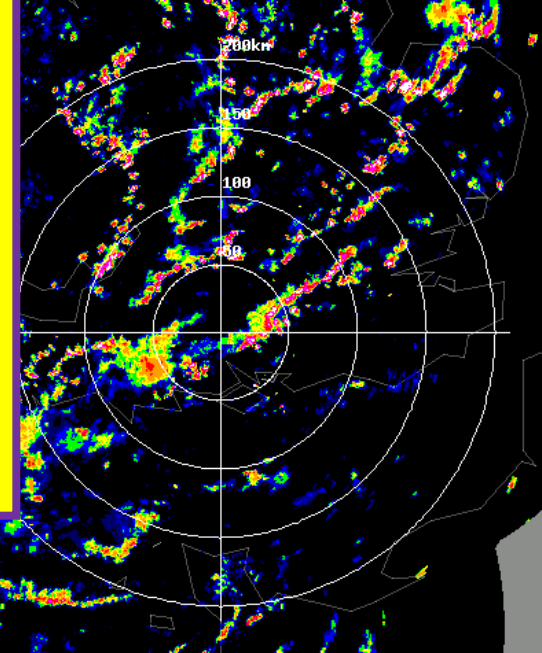
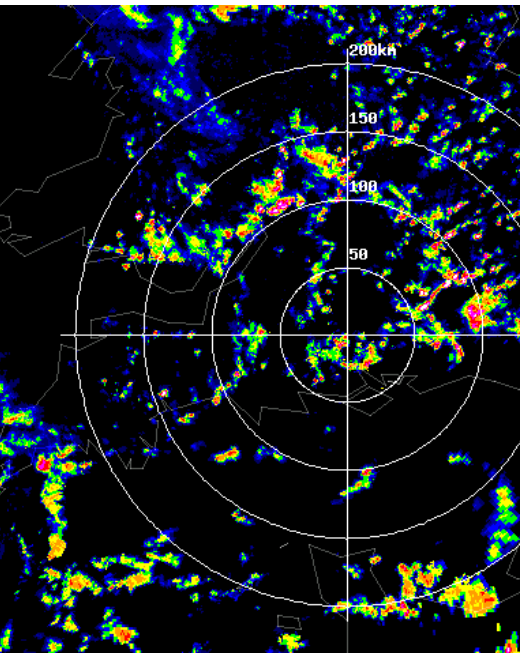


For the median life cycle (weighted by storm duration):

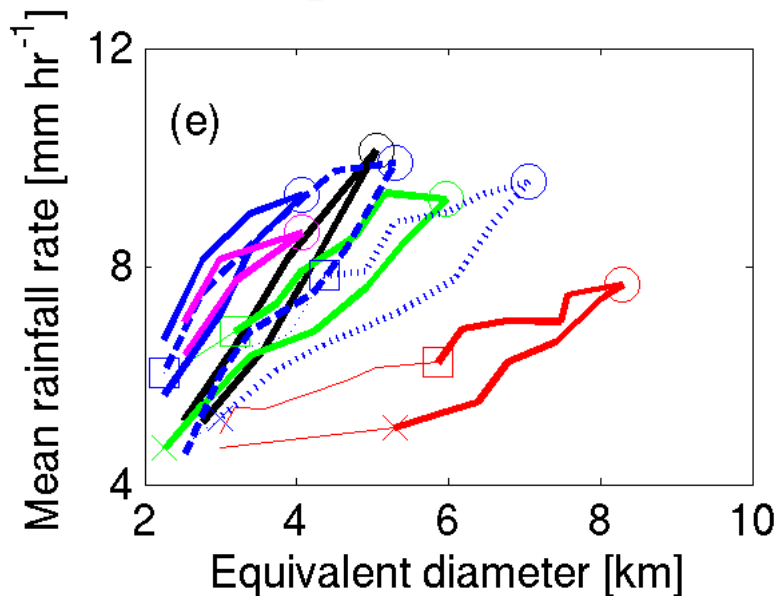
Area-integrated rainfall is reduced as grid length is reduced, with 200m and 100m models performing best.

Area-integrated rainfall is increased with increasing mixing length, with 40m (default for 200m model) best.

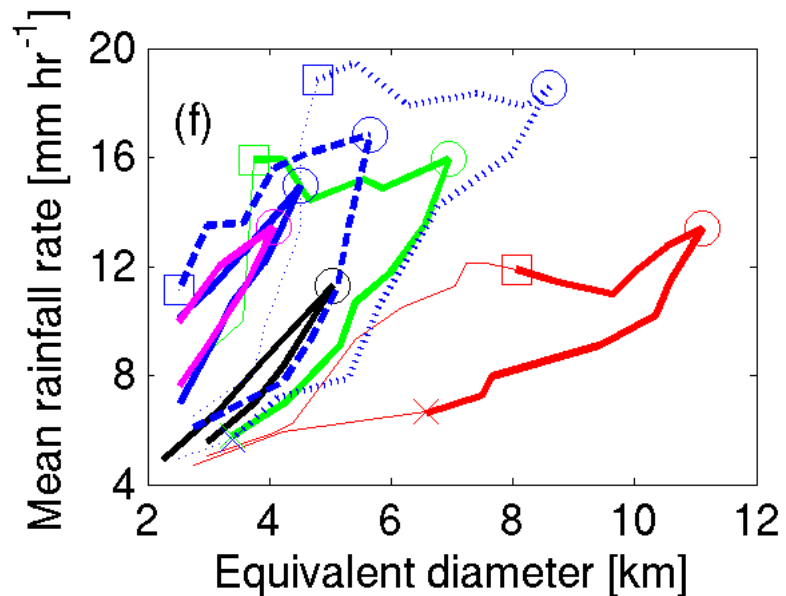
200m model,
300m mixing length



20 April 2012



25 Aug 2012

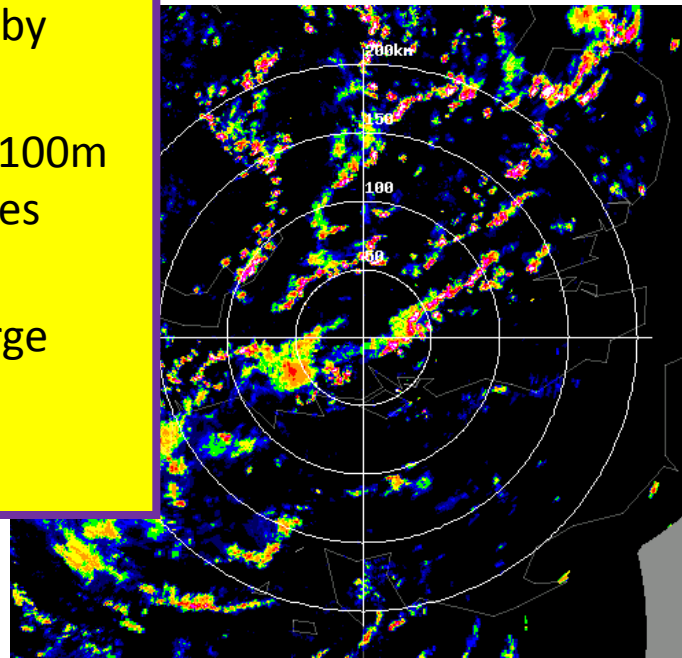
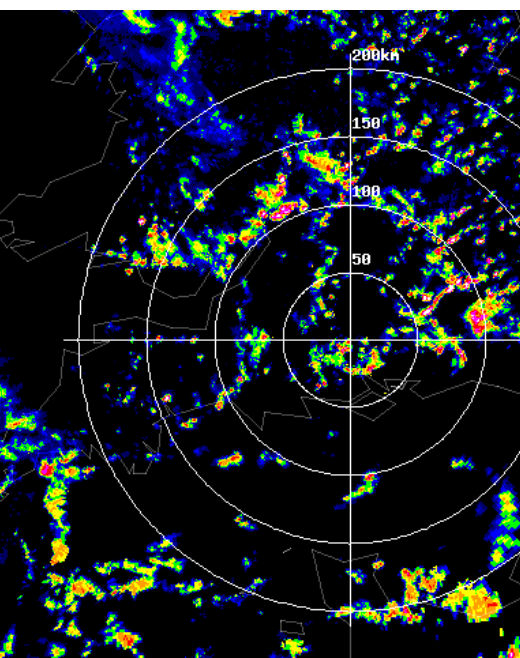


For the median life cycle (weighted by storm duration):

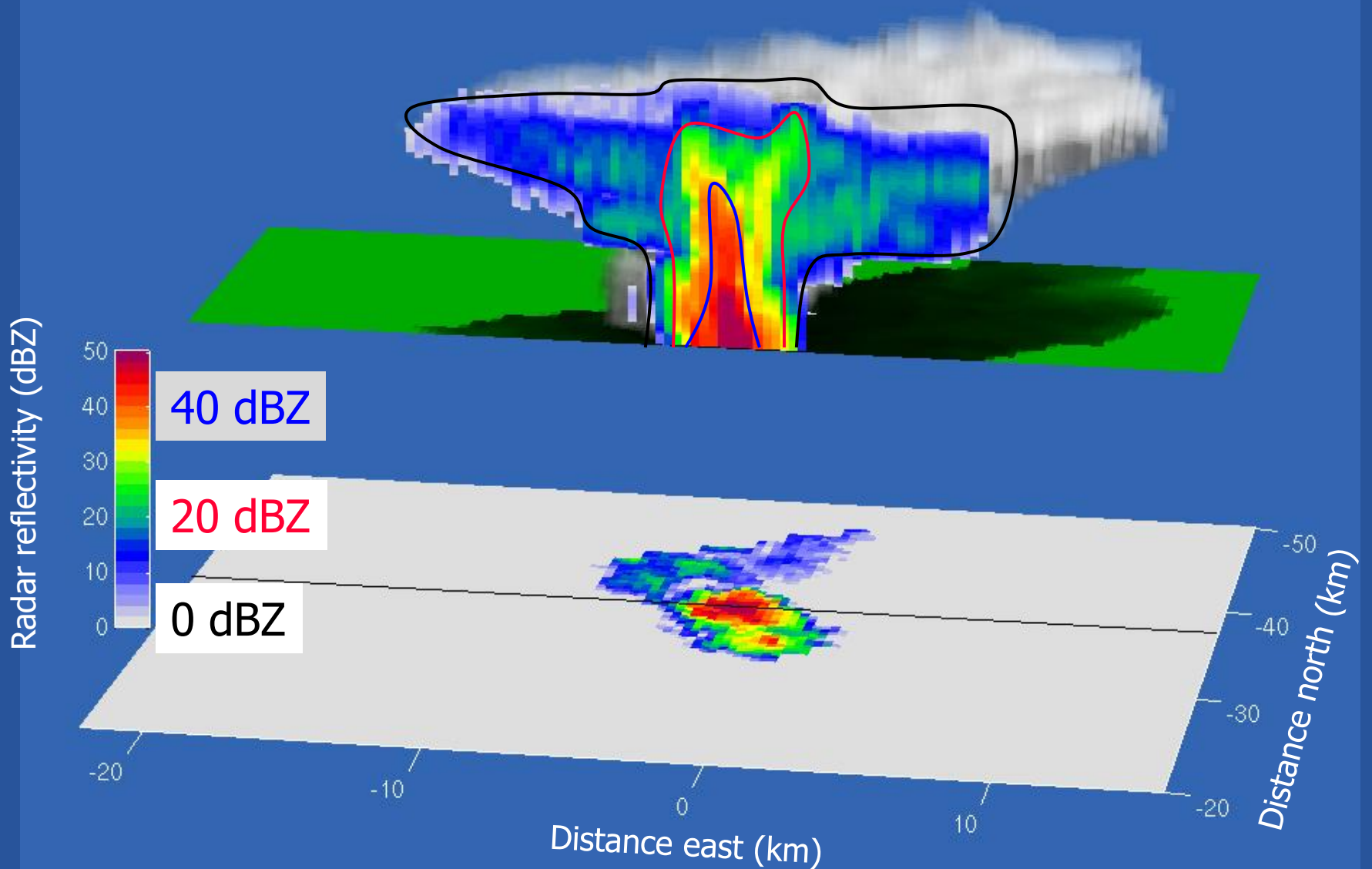
Despite good AIR cycles, 200m and 100m models have too high mean rain rates initially, also at small areas.

1500m has storms that grow too large compared to observations, though intensity is ok.

100m mixing length
200m model,
300m mixing length



Storm structure from radar



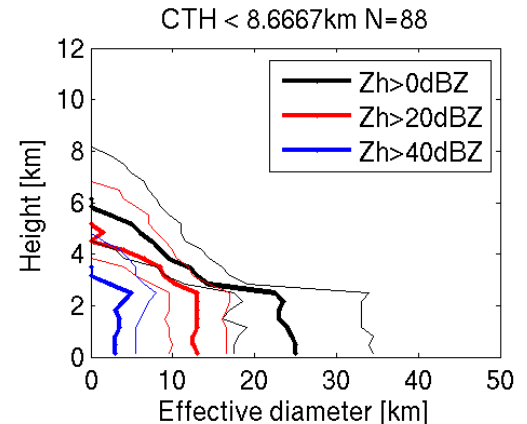
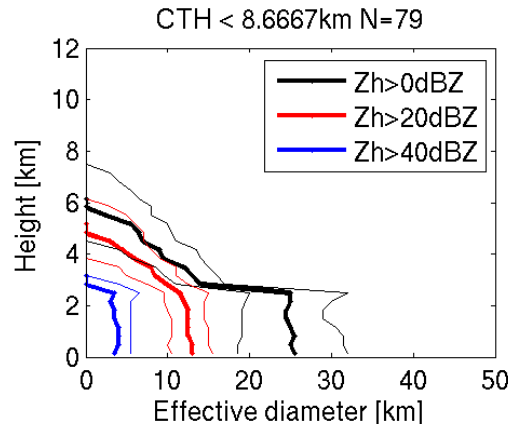
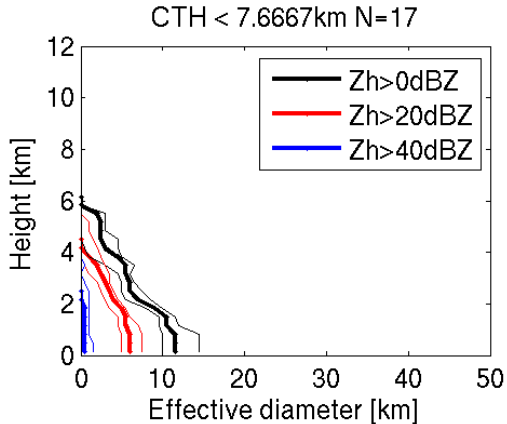
Vertical profile

Observations

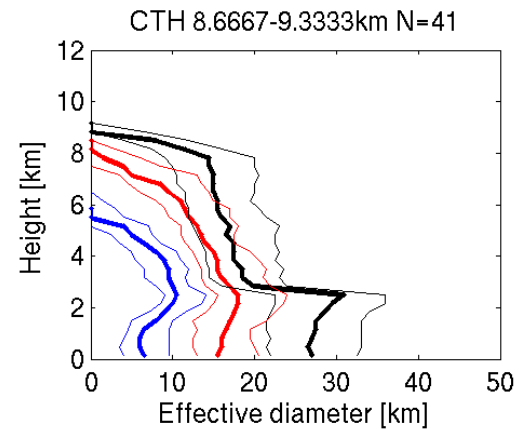
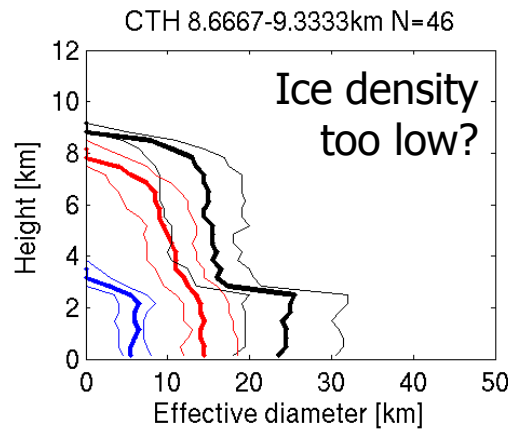
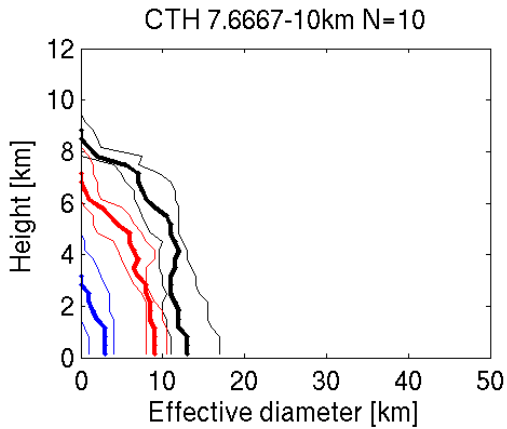
1.5-km model

1.5-km + graupel

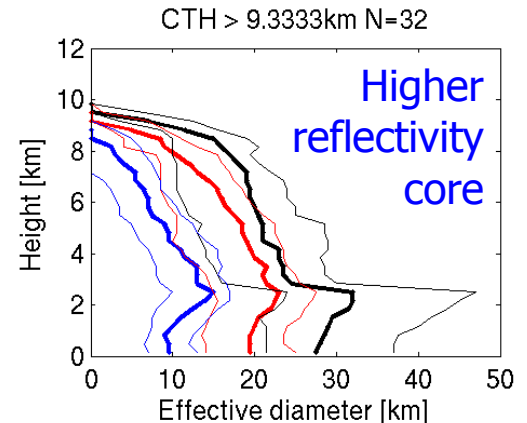
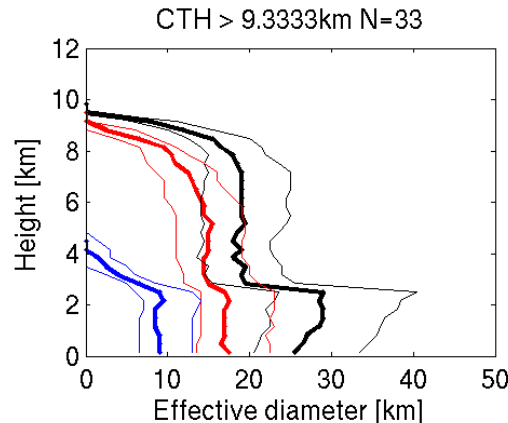
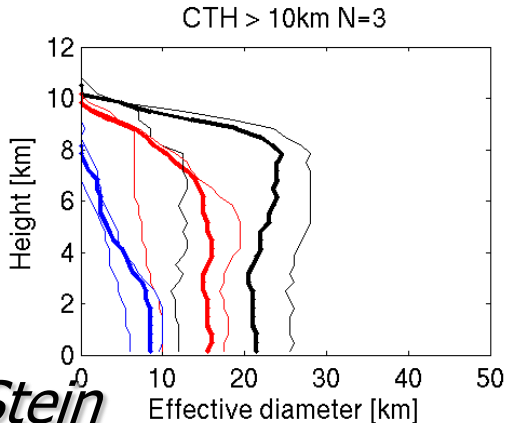
First 60% of storms by cloud-top height



Next 30%



Top 10%



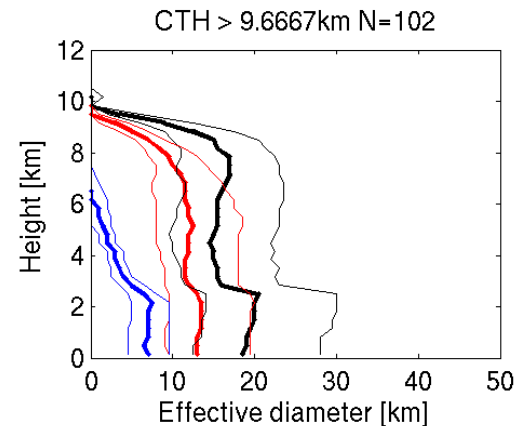
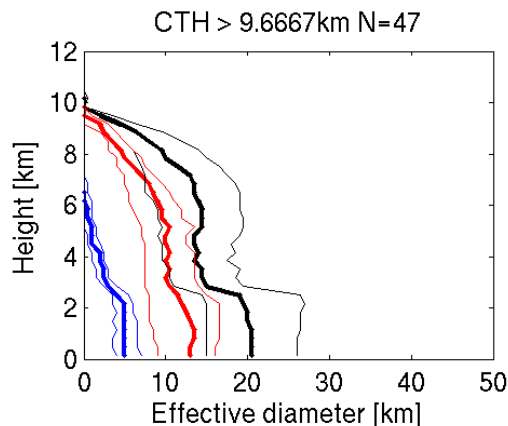
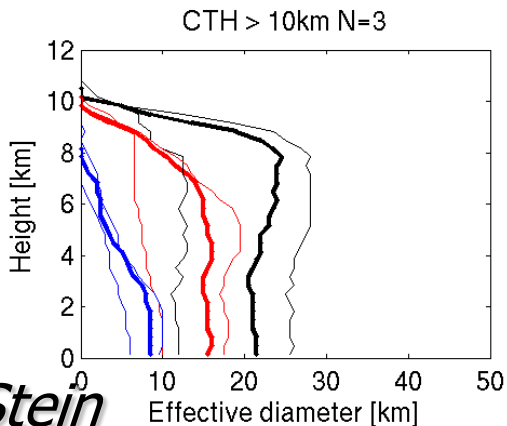
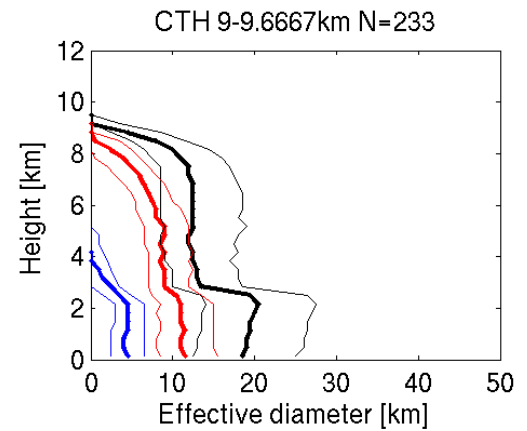
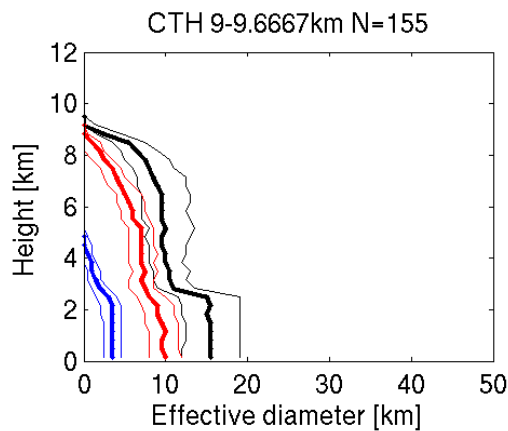
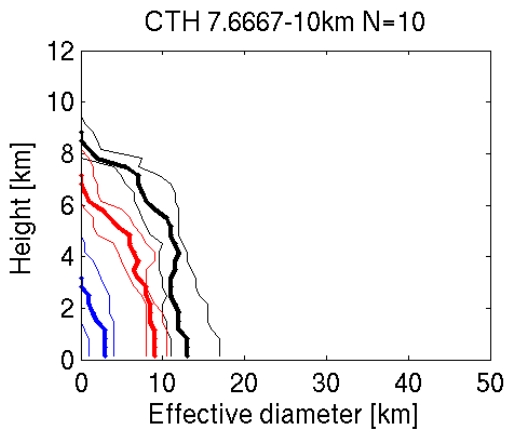
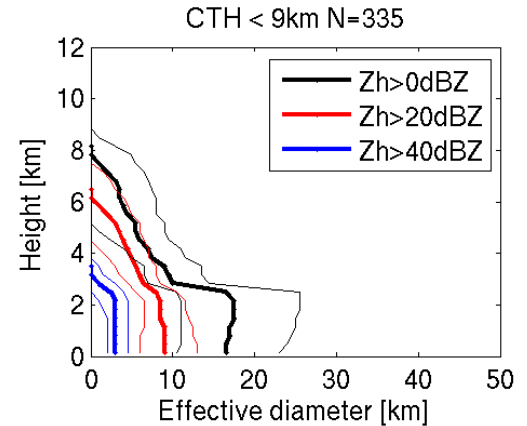
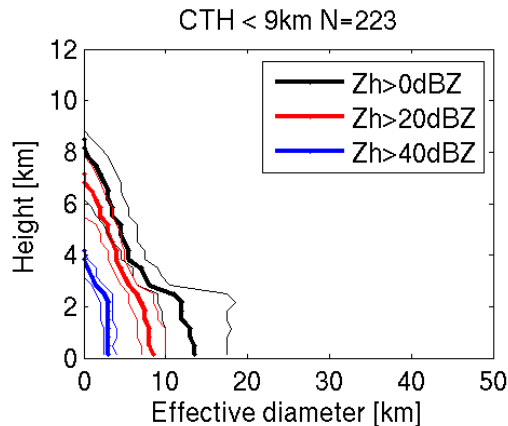
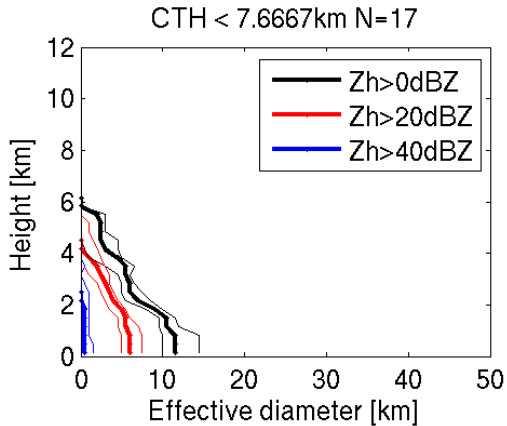
Vertical profile

First 60% of storms by cloud-top height

Observations

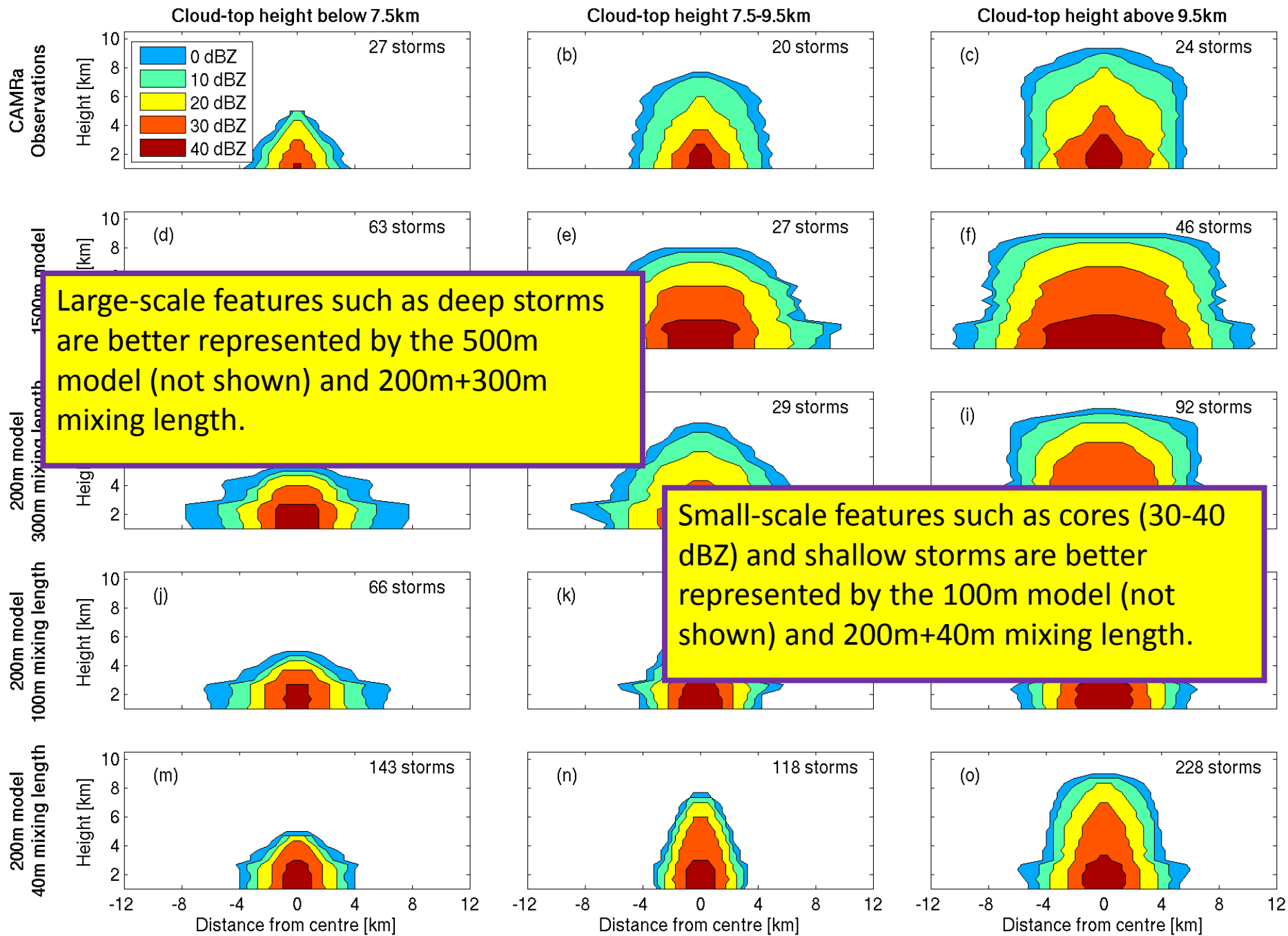
200-m model

500-m model

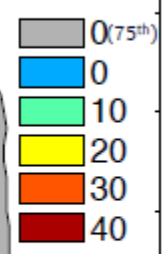
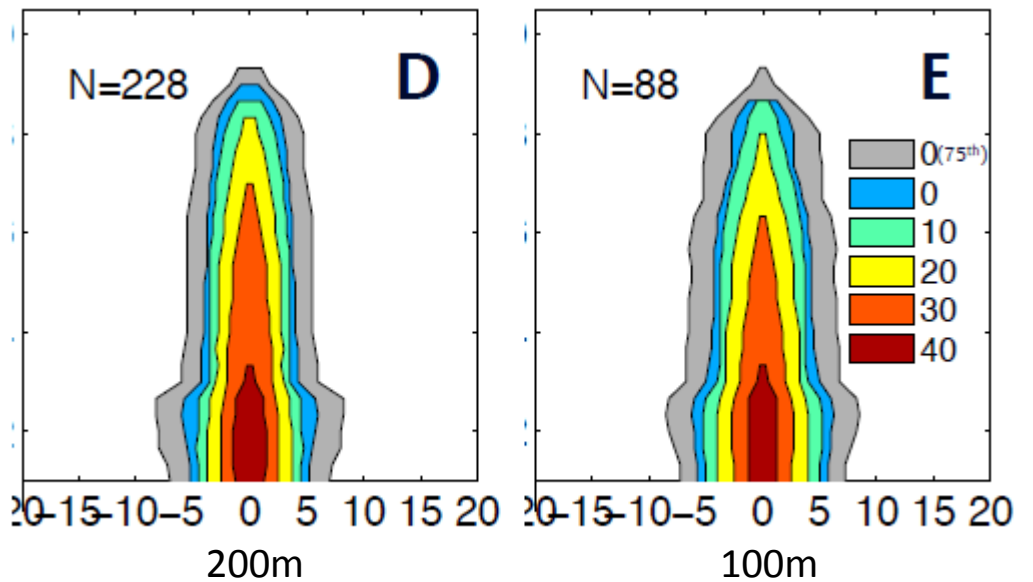
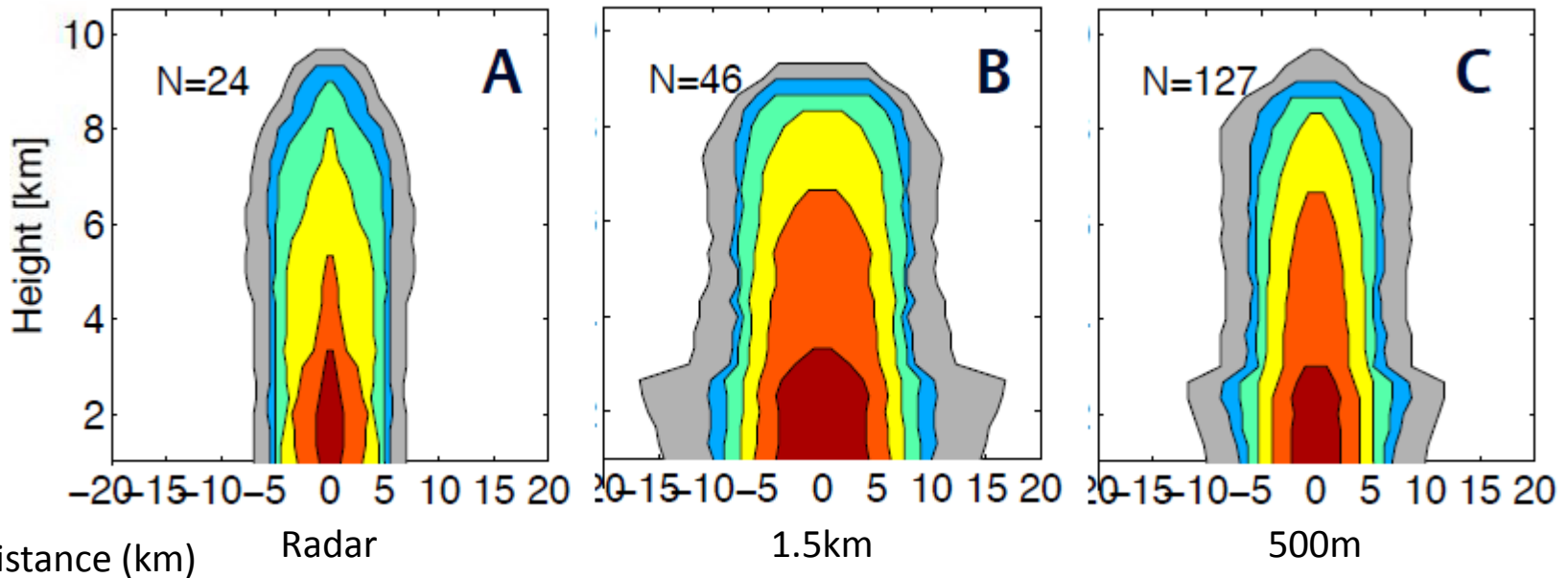


Next 30%

Top 10%



Cloud widths for different reflectivity thresholds.



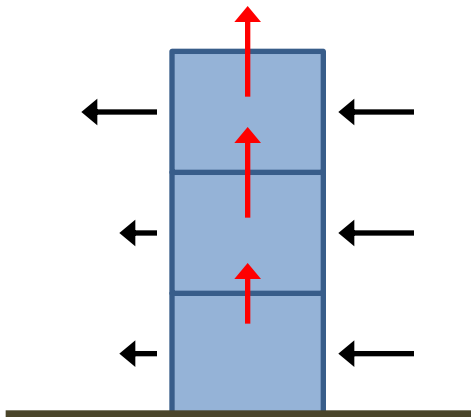
Median width of deep storms
25th Aug 2012

Thorwald Stein

Evaluation of Convective Updraughts

Estimation of vertical velocities from continuity

- Key uncertainty in models is convective updraught *intensity* and *spatial scale*
- Can we estimate updrafts from Doppler wind sufficiently well to characterize the distribution of intensity and spatial scale?



Vertical cross-sections (RHIs) are typically made at low elevations (e.g. $< 10^\circ$)

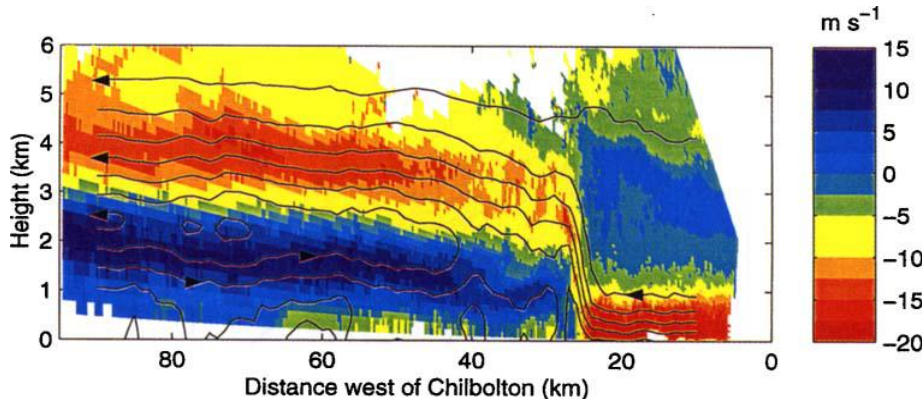
Radial velocities provide accurate estimate of the horizontal winds

Assume vertical winds are zero at the surface

Working upwards, changes in horizontal winds at a given level

increment the vertical wind up to that point

Must account for density change with height



Chapman & Browning (1998)

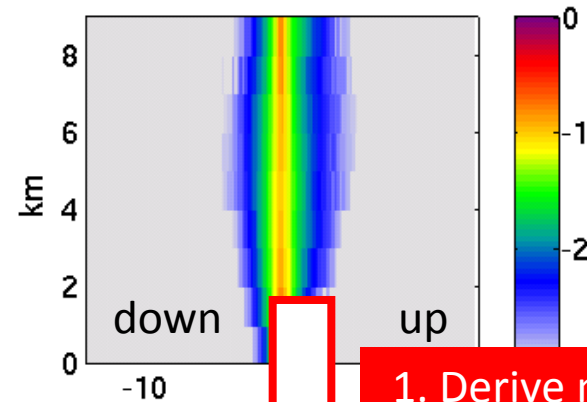
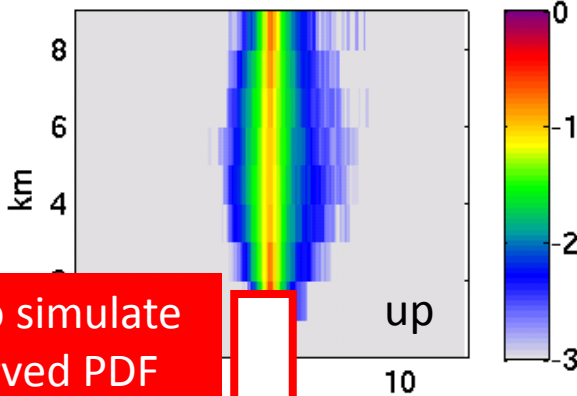
- In quasi-2D features (e.g. squall lines) can assume continuity to estimate vertical velocity

Vertical velocity distributions with height

Observations

500m

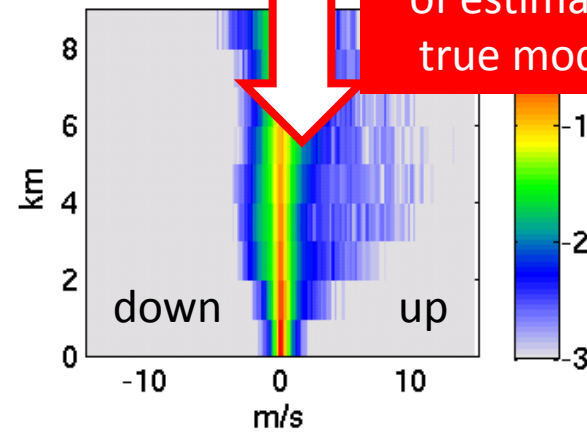
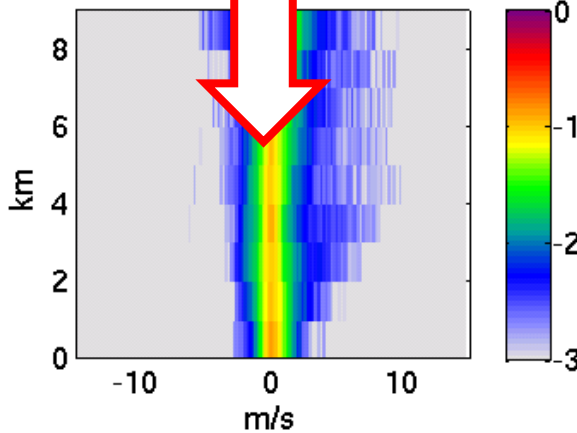
Estimated vertical velocity



2. Use map to simulate "true" observed PDF

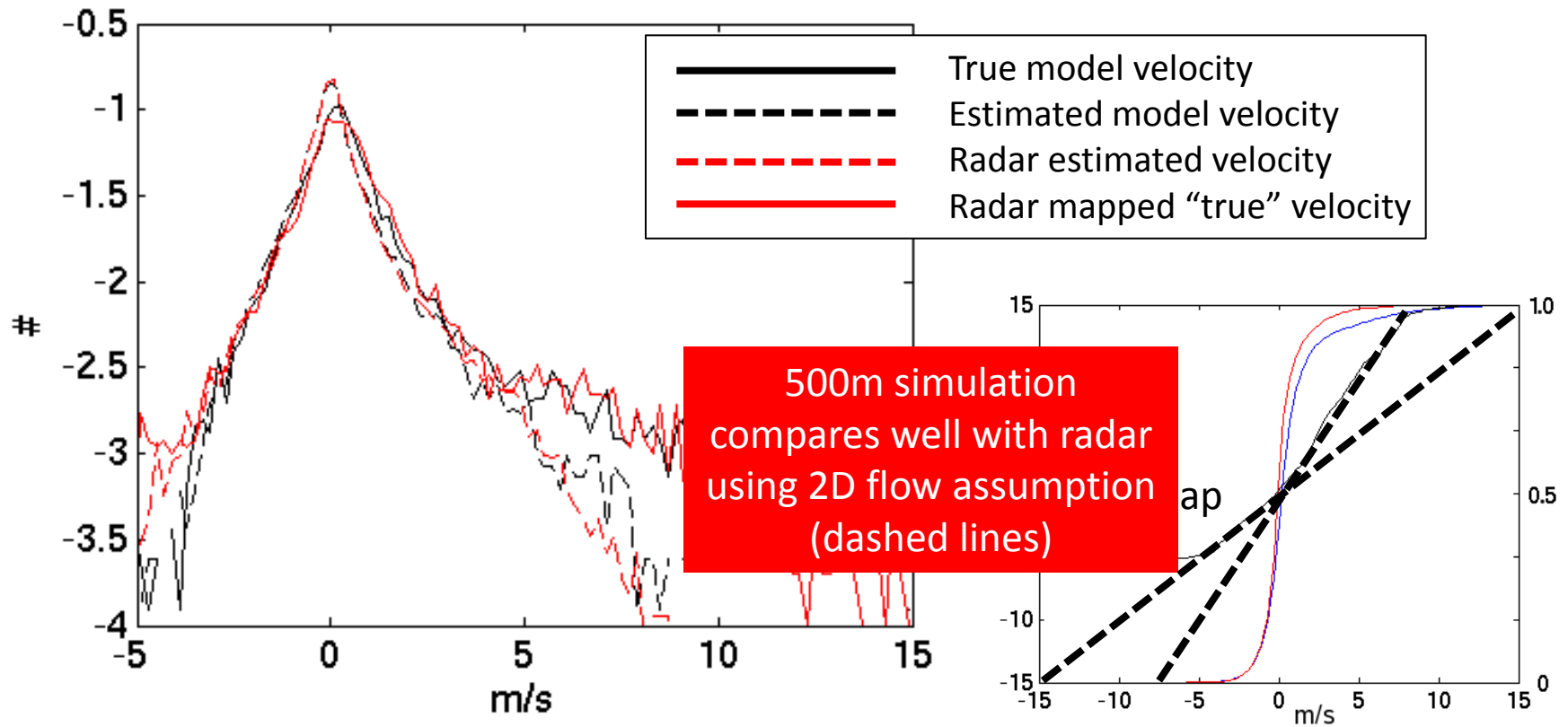
1. Derive map from PDF of estimates to PDF of true model velocities

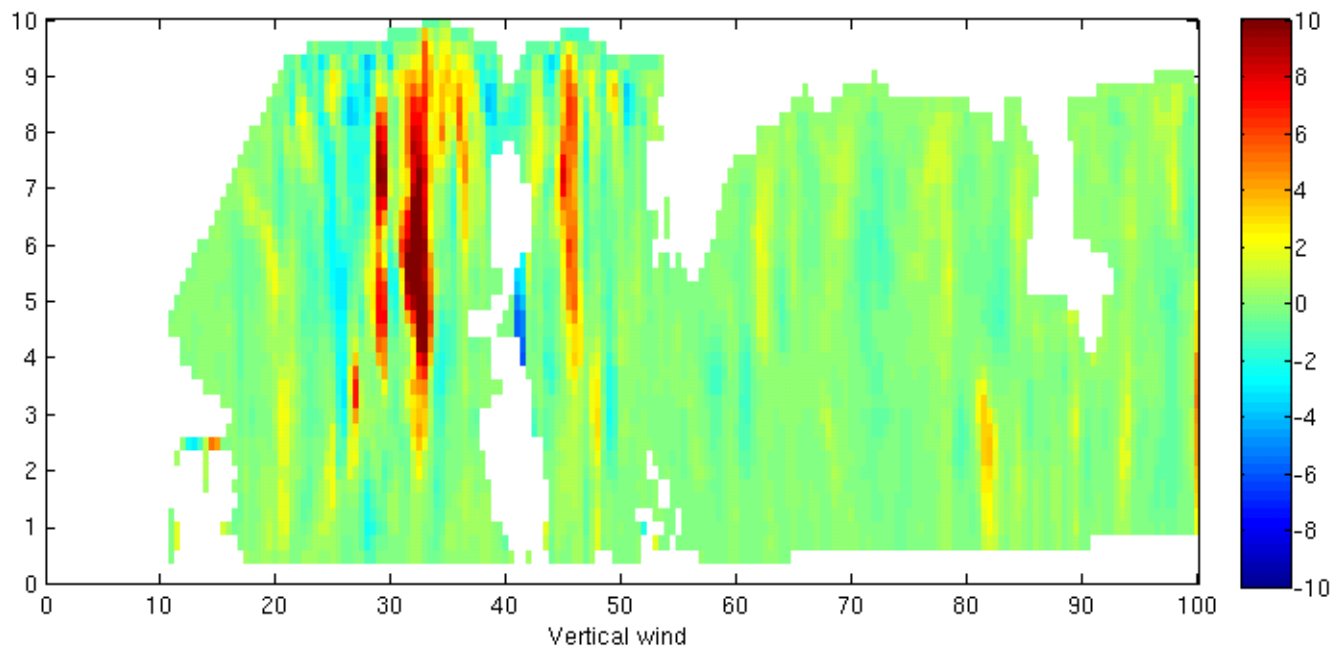
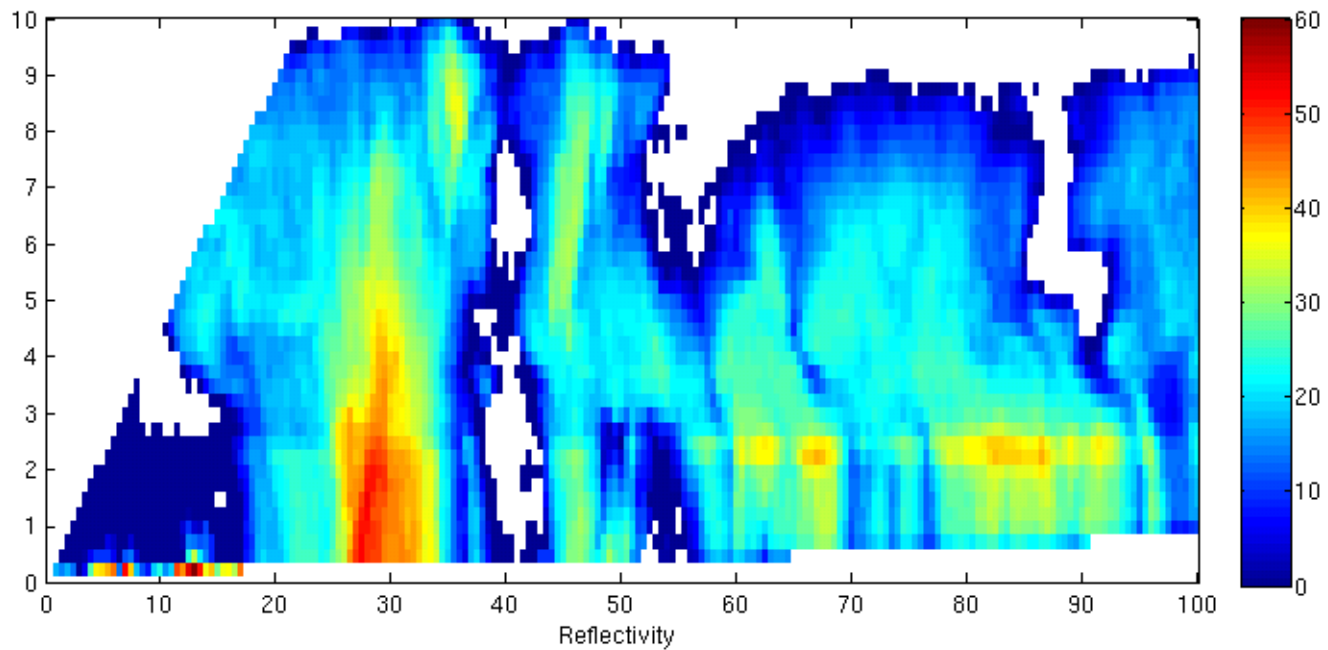
True vertical velocity



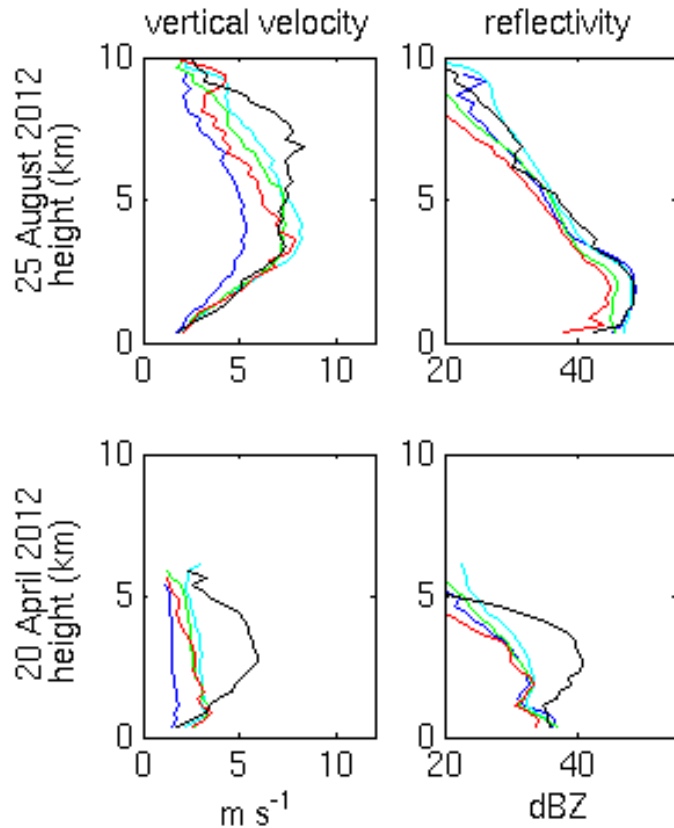
Radar data with dBZ>0 within 90 km of the radar

Vertical velocity distribution between 7-8 km





Maximum vertical velocity and reflectivity associated with each rain cell as a function of height



Rain cells defined by surface rainfall rates > 4 mm/hr over 10 km²

Only the top 50% of cells based on cell-integrated rainfall rate have been analysed for each set of simulations (so the sampling of the population of cells is similar to the cells sampled by the radar)

Two cases considered:

Deep convection (25th August 2012)

Moderate convection (20th April 2012)

Model grid length

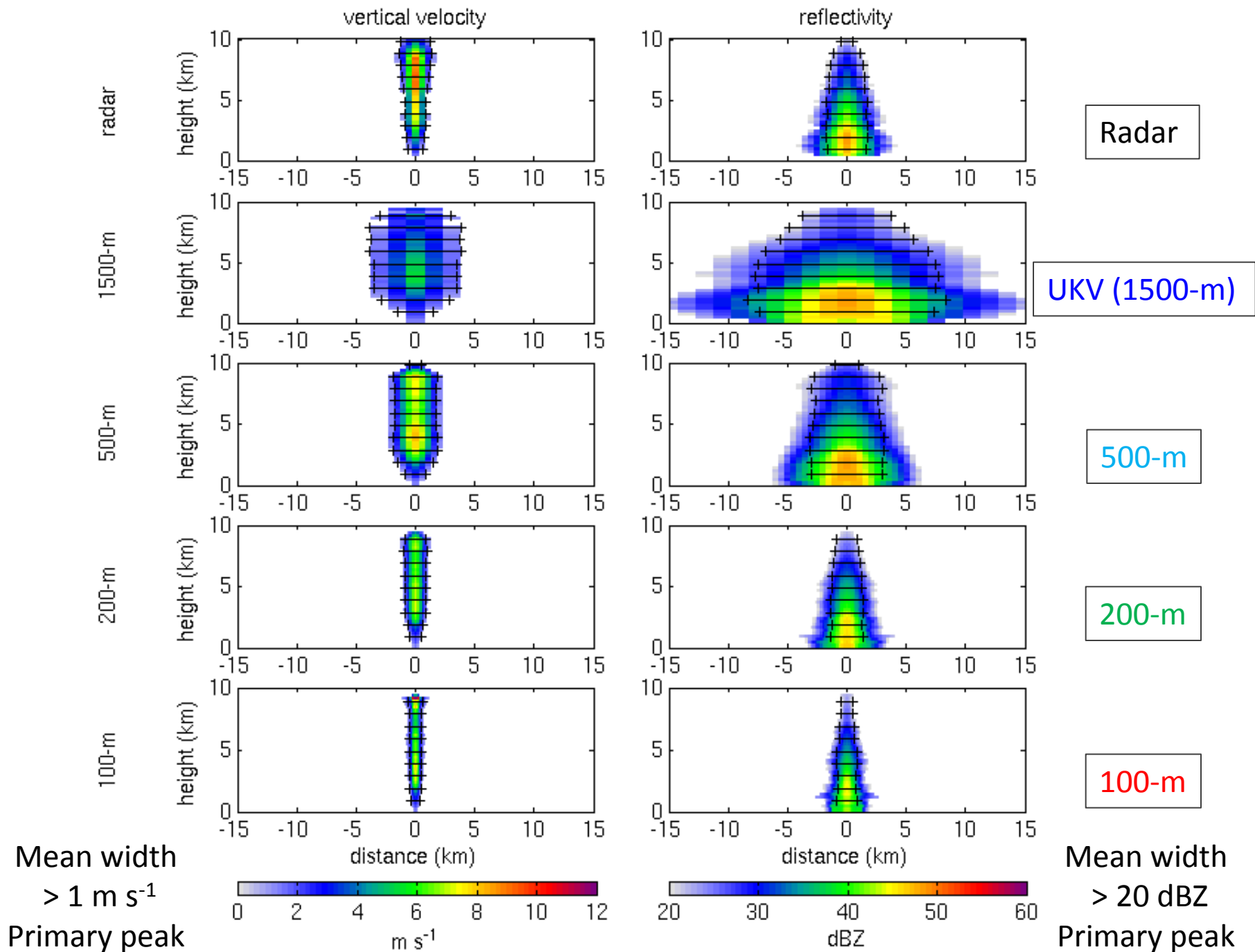
UKV (1500-m)

500-m

200-m

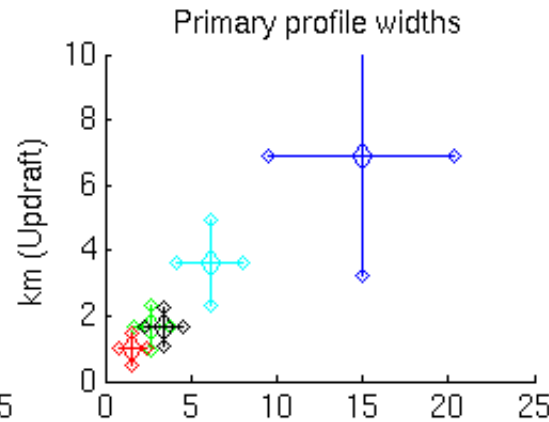
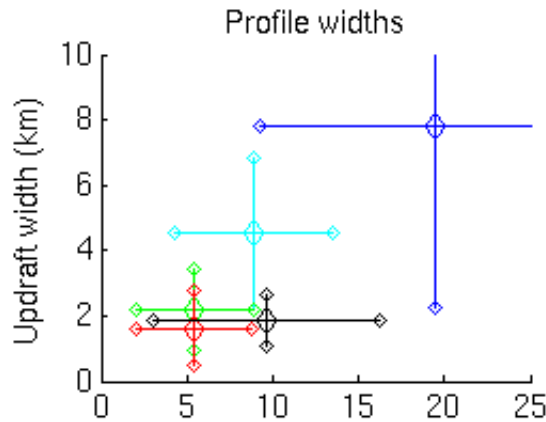
100-m

Radar

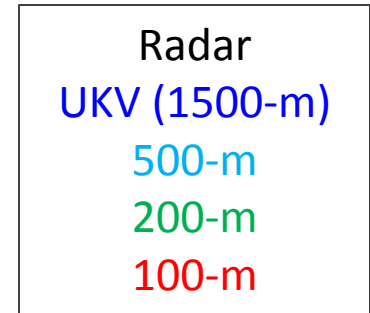


Updraught width vs. reflectivity width

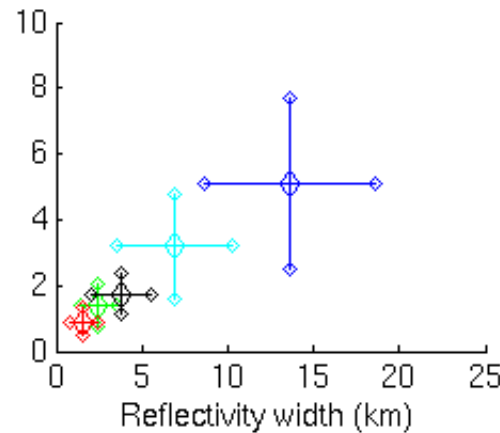
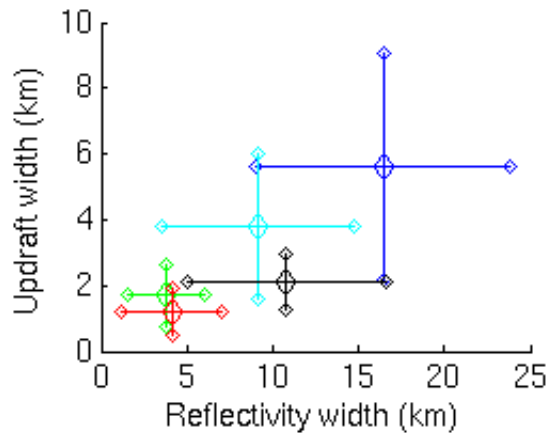
25 August 2012



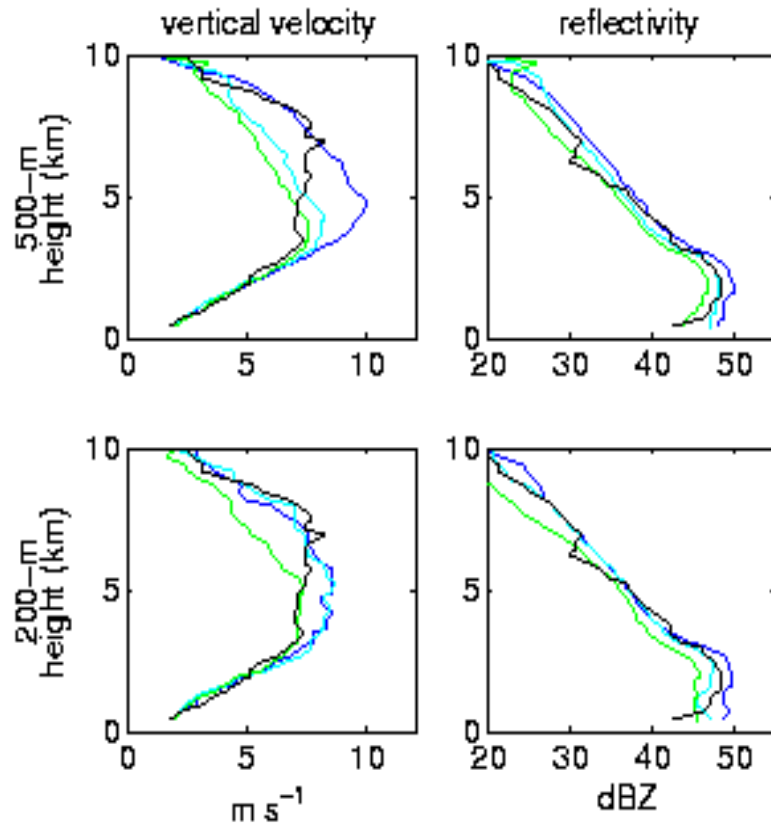
Height= 3km



20 April 2012



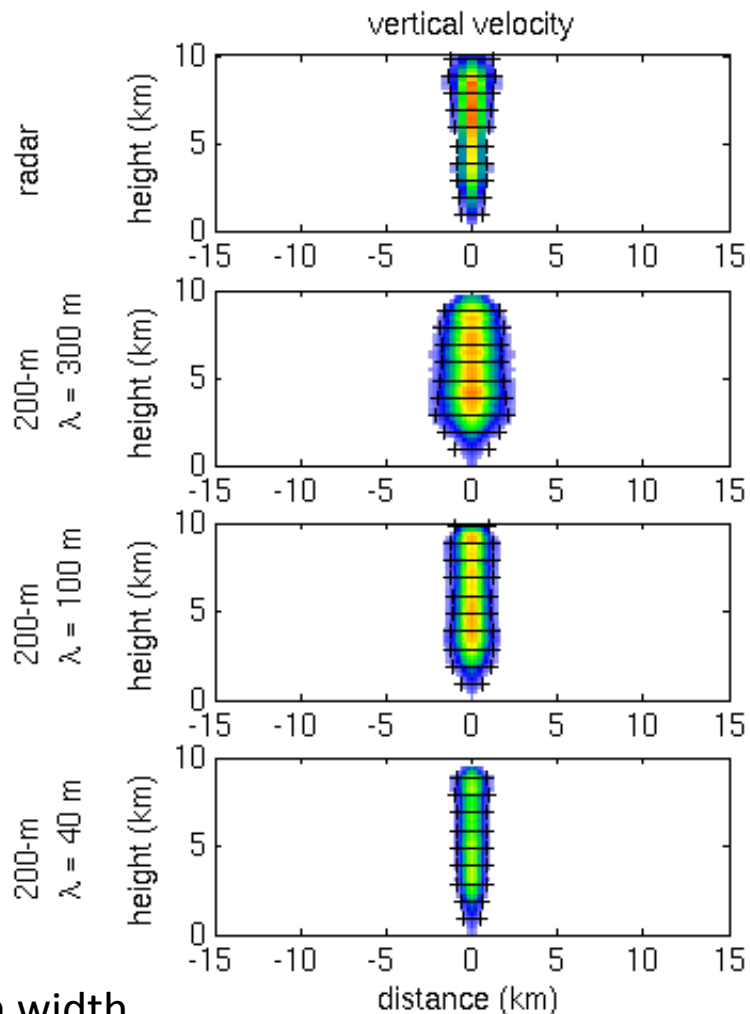
Influence of mixing length in the sub-grid turbulent mixing scheme



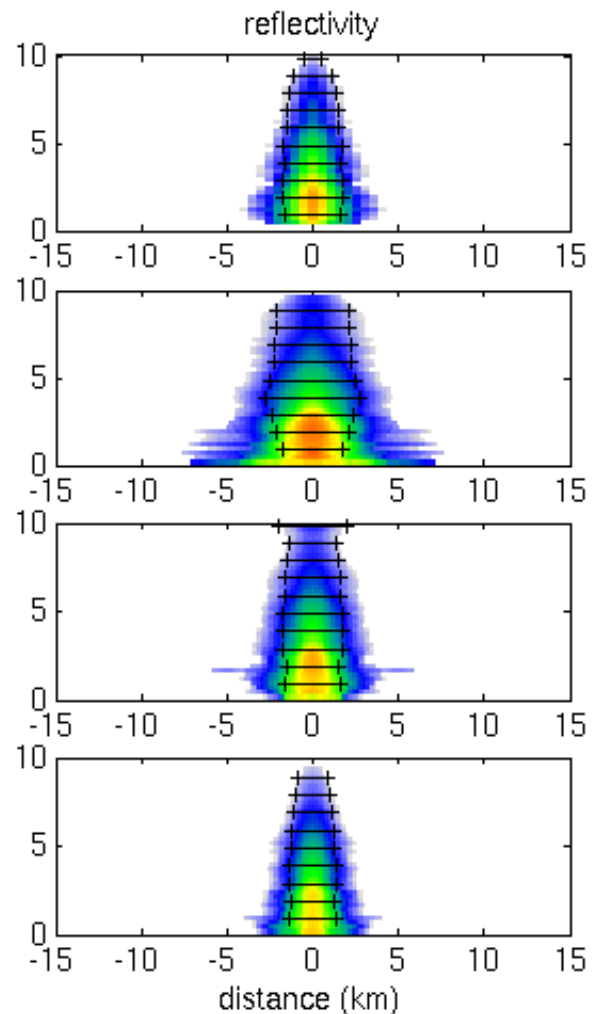
Scheme serves to diffuse
humidity, temperature and wind
fields due to unresolved
turbulence

Default configuration:
Mixing length (λ) = 20% of the
horizontal grid length

Mixing length: $\lambda=300\text{m}$ $\lambda=100\text{m}$ $\lambda=40\text{m}$



Mean width
 $> 1 \text{ m s}^{-1}$
 Primary peak

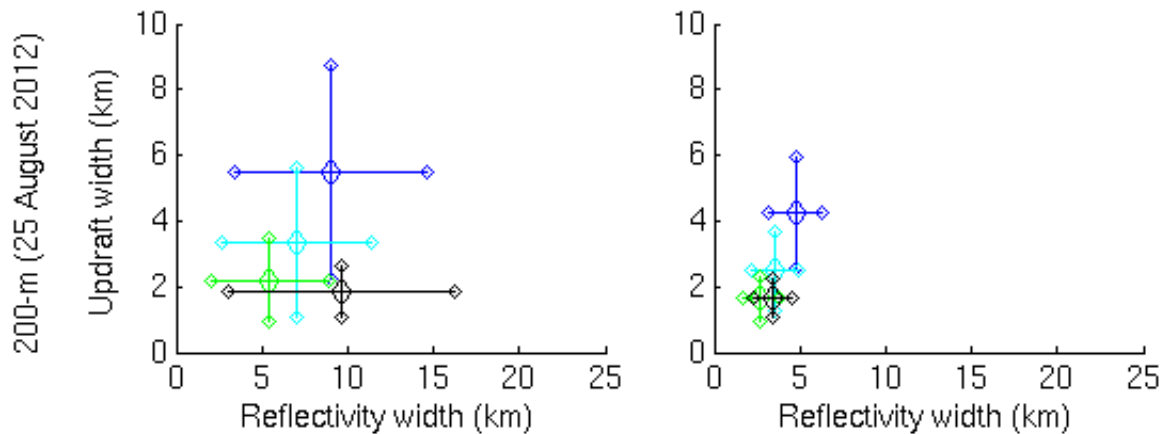
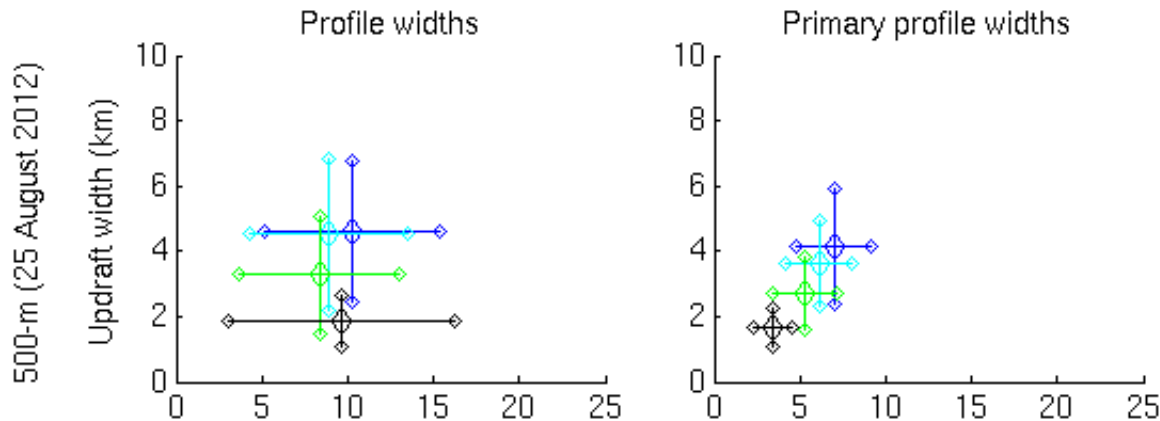


Mean width
 $> 20 \text{ dBZ}$
 Primary peak

Updraught width vs. reflectivity width

Height= 3km

500-m UM
25th August 2012



200-m UM
25th August 2012

Mixing length: $\lambda=300\text{m}$ $\lambda=100\text{m}$ $\lambda=40\text{m}$

Conclusions

- UKV under-resolves many small showers in UK while high-res models (~100m) improve some aspects but also have some problems.
- Models below 500m tend to produce too narrow showers (measured by surface rain or cloud) in cases where showers are large (for small showers 200m or 100m fits well). Cloud widths roughly the same in 200m and 100m.
- Updraught widths good in 200m model but too narrow at 100m.
- This implies that there may be an issue about how the model fills in cloud between updraft cores.
- Representation very sensitive to mixing. Also sensitivity to microphysics (fall speed).
- What can be done to understand lack of convergence and too narrow updrafts/clouds in 100m/200m models?
- Suspect problem is turbulence “grey zone”. Would better resolution of turbulence solve these problems at higher resolution? Try higher vertical/horizontal resolution and see if updraughts/clouds get wider (or stop collapsing). Work with LES community.
- Can we improve models with more appropriate subgrid mixing schemes?
- Effect of microphysics?

DYMECS

No satisfactory performance across all diagnostics

1500m model

Too long-lasting

Too much rain over lifetime

Too large rainfall area

Too large shallow structures

Too broad cores

Too broad deep structures

Weak and broad updrafts

100m model

Good duration

Good area-integrated rainfall

Too intense in early stages

Good structure for shallow

Good structure for cores

Too narrow deep structures

Good updraft strength and width

Results depend strongly on turbulent mixing length